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EXPLORING SPATIAL ASSOCIATIONS AND COMPLEX INTERACTIONS BETWEEN GEOTECHNICAL PROPERTIES AND ELECTRICAL RESISTIVITY VALUES IN CLAYEY SOILS

by

Mina Zamanian

DISSERTATION

Submitted to the Academic Faculty of The University of Texas at Arlington In Partial Fulfillment of the Requirements for the Degree of

DOCTOR OF PHILOSOPHY

The University of Texas at Arlington August 2023

EXPLORING SPATIAL ASSOCIATION AND COMPLEX INTERACTIONS BETWEEN GEOTECHNICAL PROPERTIES AND ELECTRICAL RESISTIVITY VALUES IN CLAYEY SOILS

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Dedicated to

My Mother Masoumeh, and My Father Abbas

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ABSTRACT

Exploring Spatial Association and Complex Interactions between Geotechnical Properties and Electrical Resistivity Values in Clayey Soils

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A successful design and construction of infrastructure systems such as highways and bridges highly depend on accurate estimation of geotechnical properties and understanding their spatial distributions, especially in reliability-based designs such as the load and resistance factor design (LRFD) method. Insufficient and inaccurate subsurface information has a major contribution to cost overruns and delays in up to 50% of all infrastructure projects. Insufficient site investigation may also contribute to inadequate or conservative designs, leading to costly failures or increased project costs. Hence, geophysical methods, such as electrical resistivity imaging, that can potentially transform the existing subsurface investigations are used to develop tools for subsurface characterization based on data analytic approaches. The main objective of this study is to assess the validity of the developed linear regressions in the literature by empirically evaluating one of the critical assumptions of linear regressions – independence of regression residuals. This research argues that linear regression analysis must not be used for defining the relationships between electrical resistivity and geotechnical properties since it may lead to misleading information about the subsurface conditions. First, to achieve this objective, linear regression analysis was performed on an experimental dataset to identify the impacts of geotechnical properties on electrical resistivity variations. Second, a problem was articulated with the aim of investigating whether any spatial correlation exists between geotechnical properties and electrical resistivity values. A spatial regression model was then developed that best explains the spatial variability of electrical resistivity values with the variations of geotechnical properties. The second objective of this study was to provide practical recommendations for extracting useful information from complex and non-linear interactions between geotechnical properties and electrical resistivity values with deep structures such as deep learning. The proposed approach for characterizing the soil conditions using deep learning outperformed the existing methods used in the literature.

This study identified a new research direction in the future for studying the relationships between geoelectrical and geotechnical properties through the investigation and quantification of the spatial relationships between these properties in clayey soils. The proposed approach helps create and use spatial regression models for a given site to determine the spatial distribution of geotechnical properties at each point (not necessarily those sampled using conventional site investigation methods) and conduct reliability analysis accordingly. The proposed analytical framework based on the deep learning technique also allows transportation agencies to have a better understanding of the effects of geotechnical properties on the variability of electrical resistivity values to obtain more reliable assessments of the subsurface characteristics.

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CHAPTER 1 INTRODUCTION

A successful design and construction of infrastructure systems such as highways and bridges highly depend on accurate estimation of geotechnical properties and understanding their spatial distributions, especially in reliability-based designs such as the load and resistance factor design (LRFD) method (Shahandashti et al. 2022a; Baral and Shahandashti 2022b; Sudha et al. 2009; Cosenza et al. 2006). It is also vital in identifying critical slope segments to help maintain highway embankments and cut slopes to minimize slope failures and improve transportation system efficiency (Baral et al. 2023; Baral and Shahandashti 2022a). Insufficient and inaccurate subsurface information has a major contribution to cost overruns and delays in up to 50% of all infrastructure projects (Shahandashti et al. 2022b; Baynes 2010). According to a recent nationwide study of 55 transportation agencies in the United States, the annual cost incurred due to change orders resulting from the inadequate subsurface investigation is estimated to be in the millions of dollars (Boeckmann and Loehr 2016). Insufficient site investigation may also contribute to inadequate or conservative designs, leading to costly failures or increased project costs (Adhikari et al. 2021; Shahandashti et al. 2019; Sirles 2006). Lack of continuous subsurface information may also lead to infrastructure failures caused by unforeseen circumstances (Shahandashti et al. 2021), leading to road maintenance expenses that significantly impact the state transportation budgets (Darghiasi and Shahandashti 2023a). For example, the average repair cost of karst-related damages to the infrastructures was estimated to be at least \$300 million per year in the U.S. (Weary 2015). This lack of sufficient information is due to the inherent limitation of the conventional geotechnical site investigation methods to provide continuous assessment of the subsurface. In other words, the

conventional methods only sample and provide information about a small percentage of a total sample space (Shahandashti et al. 2021). Meanwhile, the Federal Highway Administration (FHWA) has identified several subsurface exploration technologies through the EDC-5 program that are proven effective in evaluating geological, hydrological, geotechnical, and environmental site assessments. Despite the evident advantages of these technologies that can potentially transform existing subsurface investigations, many of these technologies are underutilized by many state departments of transportation because of a lack of proven implementation details for different applications, geotechnical conditions, and operational environments (FHWA 2018; Rosenblad and Boeckmann 2020). These methods offer a unique opportunity to mitigate repairing costs and limitations of conventional geotechnical site investigation methods by providing a rapid and continuous assessment of subsurface conditions using a non-invasive, and cost-effective method (Zamanian et al. 2023b).

Among the geophysical methods, the Electrical Resistivity Imaging (ERI) technique is widely used in the literature to characterize the geotechnical properties in clayey soils based on the electrical resistivity values. Empirical and analytical studies have been conducted to establish statistical models using linear, power, and exponential regression functions to investigate the effects of different hydraulic and solid phase properties of clayey soils on electrical resistivity values. Although most of these studies presented models with a relatively high goodness-of-fit, none of them have investigated the spatial association between the electrical resistivity values and geotechnical properties. The presence of autocorrelated residuals in the standard regression model leads to wrong interpretations of the regression parameters and goodness-of-fit of the models. Spatial regression models consider the spatial dependence of the error terms to accurately determine the effects of a change in geotechnical properties on the variability of electrical resistivity values. Yet there is a lack of an analytical tool for exploring the complex and non-linear relationship between electrical resistivities and geotechnical properties.

The goals of this research are: (1) to explore and quantify the spatial association between geotechnical properties and electrical resistivity values by spatial regression analysis and propose the most appropriate spatial regression model that best explains the variability of electrical resistivity values with the variations in geotechnical properties, and (2) to propose an analytical approach for extracting meaningful information from complex and non-linear interactions between geotechnical properties and electrical resistivity values by deep learning model to overcome the limitations of the linear regression analysis.

Chapter 2 presents a comprehensive literature review on the influencing geotechnical properties affecting electrical resistivity values and the existing empirical correlations between them in the literature. Chapter 2 also describes the gaps in knowledge and research objectives. Chapter 3 elaborates on a methodology for assessing the spatial autocorrelation in the regression residuals and developing an appropriate spatial regression model that best explains the variability of electrical resistivity values with the variations in geotechnical properties. Chapter 4 describes a methodology for developing a deep learning model to explore the non-linear and complex relationship between electrical resistivity values and geotechnical properties. Chapter 5 presents the conclusion.

3

CHAPTER 2 BACKGROUND

2.1. Electrical Resistivity Imaging Technology

Electrical Resistivity Imaging (ERI) technology employs fundamental physics principles of Ohm's law to determine the resistance of soil, rock, and groundwater to the flow of electrical current (Kearey et al. 2013). The ERI technology is used to uncover the horizontal and vertical discontinuities in the earth's materials. The soil electrical resistivity is a function of soil and rock matrix, degree of saturation, pore fluid conductivity, soil fabric structure, and soil compressibility (Ekwue and Bartholomew 2010; Samouëlian et al. 2005; Lapenna et al. 2005; Friedman 2005; Giao et al. 2003; Rinaldi and Cuestas 2002; Yang 2002). The soil's electrical properties can be studied by inducing a direct or a very low-frequency current into the ground. The current is induced into the ground across two electrodes (current electrodes), and then the resulting voltage is received by the other two electrodes (potential electrodes) (ASTM Standard D6431-18 2018; ASTM Standard D6429-99 2011). In practice, a large number of electrodes (e.g., 28, 56, or more) and multi-electrode cables are used to speed up the data acquisition and improve the quality of large datasets (Akingboye and Ogunyele 2019). Simultaneous measurements can be recorded using a multi-electrode array; a switching box automatically selects and switches the relevant four electrodes based on the predefined sequence stored in the resistivity meter (Bernard et al. 2006).

The main benefit of electrical resistivity imaging over the other advanced geophysical methods is its wide range of applications in determining various subsurface anomalies and soil properties. Figure 2.1 shows a comparison of the number of applications of advanced geophysical

tools in the subsurface investigation (ASTM Standard D7400-19 2019; ASTM Standard D5753-18 2018; ASTM Standard 6285-99 2016; Rivers 2016; Li et al. 2014; ASTM Standard D5778-12 2012; ASTM Standard D6429-99 2011; British Standards Institution 2010; Edet 2009; Rogers 2009; Anderson et al. 2008; Sirles 2006; Fenning and Donnelly 2004; Wightman et al. 2004; Williams and Johnson 2004).



Notes: "ERI" denotes Electrical Resistivity Imaging, "S.Refr." denotes Seismic Refraction, "S.Refl." denotes Seismic Reflection, "MASW" denotes Multi-channel Analysis of Surface Wave, "IP" denotes Induced Polarization, "MWD" denotes Measurements While Drilling, "SASW" denotes Spectral Analysis of Surface Wave, "SP" denotes Self Potential, and "SCPT" denotes Seismic Cone Penetration Test.

Figure 2.1 A comparison of the number of applications of advanced geophysical methods

(Source: Adapted from Shahandashti et al. 2021)

The capital cost of ERI equipment is evaluated at around \$60,000 (AGI, IRIS Instrument, GuideLineGeo) and a total annual salary of \$210,000 is considered for a crew of three persons to perform the ERI surveys and data analysis. In addition to providing a continuous assessment of subsurface conditions, the ERI incurs no additional costs. On the other hand, aside from the equipment cost of conventional geotechnical site investigation such as CPT and SPT which starts from \$150,000 (TMG Manufacturing), the soil test drilling incurs a cost of \$156 and \$77 per meter of advancement through the depth, respectively (Crisp et al. 2018).

2.2. Electrical Mixing Model

Electrical mixing models describe how the bulk electrical resistivity of a conducting medium is associated with the resistivity of components of porous media. According to Archie (1942), the bulk electrical resistivity of fully saturated coarse-grained soils is related to the geometry of pore spaces and pore fluid's electrical resistivity. Keller and Frischknecht (1966) later expanded Archie's model for partially saturated porous media as given by Equation 2.1.

$$\rho = a\rho_w n^{-m} S^{-p}$$
 Eq. 2.1

where ρ is bulk electrical resistivity, ρ_w is pore water resistivity, *a* is compaction constant, *n* is porosity, *S* is degree of saturation, *p* is saturation parameter, and *m* is cementation parameter. The cementation parameter depends on the pore tortuosity and pore network interconnectivity, and the saturation parameter represents the pore water in the soil matrix. The values of *a*, *p*, and *m* are typically obtained by regression analyses (Bryson 2005).

A generalized form of Archie's model for fine-grained soil that includes the effect of surface conductivity in the cementation factor is as follows (Shah and Singh 2005):

$$\sigma_b = c\sigma_w \theta^m$$
 Eq. 2.2

where σ_b is bulk electrical conductivity, *c* is a fitting parameter, σ_w is pore water conductivity, θ is volumetric water content, and *m* is cementation parameter. The values of *c* and *m* are calculated by Equations 2.3 and 2.4 for soils with a clay fraction above 5%.

$$c = 0.6 \ CL^{0.55}$$
 Eq. 2.3

$$m = 0.92 \ CL^{0.2}$$
 Eq. 2.4

where CL is the percentage of clay. For soils with less than 5% clay fractions, values of 1.45 and 1.25 are considered for c and m. In clayey soils, the electrical current flows through pore space by the movement of ions in pore water and surface charges at the soil and water interface (Rhoades et al. 1989). Therefore, the specific surface area and surface conductance of clayey soil particles, which also correlate with the residual friction angle (Tiwari and Marui 2005), affect the electrical resistivity (Klein and Santamarina 2003).

2.3. Existing Relationships between Geotechnical Properties and Electrical Resistivity Values

2.3.1. Regression Analysis

The regression analysis has a variety of applications in assessing construction and transportation infrastructure resiliency (Zamanian et al. 2024; Darghiasi et al. 2023b and 2023c; Zamanian and Shahandashti 2022). Likewise, in the field of geotechnical engineering, various empirical and analytical studies have been conducted to develop statistical models using linear, power, and exponential regression functions to investigate the effects of different hydraulic and solid phase properties of clayey soils on electrical resistivity. Among the hydraulic properties, soil water content has been identified as one of the significant factors affecting soil electrical resistivity (Zamanian and Shahandashti 2022; Robinson et al. 2008; Samouelian et al. 2005; Friedman 2005). Besson et al. (2010) also showed that 48% of the total variations of the electrical resistivity are attributed to the volumetric water content of the soil. The soil's electrical resistivity decreases as the water content increases since the electrical current is transmitted through the movement of ions in pore water (Siddiqui and Osman 2012). The indirect relationship between the water content and electrical resistivity of clayey soils was also identified by Shahandashti et al. (2021), Rezaei et al. (2018), Abidin et al. (2013), Siddiqui and Osman (2012), Kibria and Hossian (2012), and Michot et al. (2003). Abu-Hassanein et al. (1996) investigated the effect of degree of saturation on the

soil's electrical resistivity. They concluded that an increase in the degree of saturation of clayey soil leads to a decrease in the electrical resistivity values. Rinaldi and Cuestas (2002) showed that the void ratio (one of the controlling factors of gravimetric water content) significantly affects electrical resistivity variations. The electrical resistivity of clayey soil decreases as the dry unit weight increases while keeping gravimetric water content constant (Lin et al. 2016). Nevertheless, the variability of electrical resistivity is less sensitive to the variations of dry unit weight than the gravimetric water content, and it is almost negligible at the gravimetric water contents above 30% (Kibria and Hossain 2012). Rashid et al. (2018) observed a 50% reduction in the electrical resistivity for a 20% increase in the dry density. An increase in dry density results in less pore space and more interparticle contacts, decreasing soil resistance to electrical current flow. The rate of reduction in the electrical resistivity with increasing dry density depends on soil type. In another study, Alsharari et al. (2020) assessed the combined effects of gravimetric water content, dry unit weight, salinity, and percentage of a clay mineral on the variability of electrical resistivity of clayey soils using multiple regression analysis. The effects of Atterberg limits on the variations of electrical resistivity were studied by Abu-Hassanein et al. (1996) and Long et al. (2012). They showed that the lower electrical resistivity values are associated with the higher plasticity index/liquid limit measures. Lin et al. (2016) also showed that the electrical resistivity of clayey soils is more correlated to the plasticity index than the liquid limit. Abu-Hassanein et al. (1996) also concluded that the percentage of fines (percent of soil finer than 75 microns) or percentage of clay (percent of soil finer than 2 microns) of soils impacts the electrical resistivity of fine-grained soils. Soils with more percentage of fines and clays yield lower electrical resistivity values because they have higher specific surface areas, which promotes the transmission of electrical current (Morin 2006).

Tables 2.1 to 2.3 show examples of empirical studies relating the electrical resistivity value to the geotechnical properties. The standard linear regression model has been widely used to explain the variability of electrical resistivity with the water content, plasticity index, void ratio, and porosity. The second-order regression models (quadratic regression models) are proposed by Lin et al. (2016) and Kibria and Hossain (2012) to study the effects of unit weight on the variations of electrical resistivity values. The power law and exponential regression functions have also been used to provide estimates for the unit weight, degree of saturation, and porosity using electrical resistivity values.

Authors	Soil type	No. of data points	Correlation	Parameter values ^a	Coefficient of determination
Goyal et al.			Linear, w - ρ	<i>a</i> = 500, <i>b</i> = -10	0.980
1996					
Michot et al.	Loamy	30-250	Linear, <i>w</i> - <i>p</i>	<i>a</i> = 28.5 to 37.7,	0.212 - 0.941
2003	clay			b = -0.05 to 0.36	
Cosenza et al.	Sand	20	Power law, <i>p</i> -w	<i>a</i> = 1.187, <i>b</i> = -2.444	0.821
2006	and clay				
Fallahsafari et	Clay	25	Exponential, w - ρ	<i>a</i> = 21.66, <i>b</i> = -0.19	0.619
al. 2010			Exponential, γ_d - ρ	<i>a</i> = 11426, <i>b</i> = 0.181	0.568
			Linear, e -ln(ρ)	<i>a</i> = 0.702, <i>b</i> = -0.36	0.484
			Linear, $n-\ln(\rho)$	<i>a</i> = 0.415, <i>b</i> = -0.18	0.480

 Table 2.1 Examples of empirical studies relating electrical resistivity to the geotechnical properties (before 2010)

Notes: " ρ " denotes electrical resistivity, "w" denotes water content, " γ " denotes bulk unit weight, " γ_d " denotes dry unit weight, "PI" denotes plasticity index, "e" denotes void ratio, "n" denotes porosity, and " S_r " denotes degree of saturation.

^{*a*} Coefficient of *a*, *b*, and *c* represent constant parameters in the linear (y=a+b.x), power law ($y=a.x^b$), exponential (y=a.exp(b.x)), and quadratic ($y=a.x^2+b.x+c$) regression functions.

Authors	Soil type	No. of data points	Correlation	Parameter values ^a	Coefficient of determination
Kibria and Hossain 2012	Clay	59	Linear, ρ -w	a = 119.26 to 328.03,	0.810 - 0.880
110354111 2012				<i>b</i> = -1.094 to -1.351	
			Power law, ρ - S_r	<i>a</i> = 2.41 to 2.73,	0.550 - 0.960
				<i>b</i> = -1.64 to -0.58	
			Quadratic, ρ - γ	a = 0.095 to 0.7107,	0.98 - 1.0
				b = -24.541 to 3.461,	
				<i>c</i> = 34.099 to 217.98	
Siddiqui and Osman 2012			Linear, $w-\ln(\rho)$	<i>a</i> = 0.644, <i>b</i> = -0.0451	0.659
0000000000000			Power law, γ - ρ $a = 14.9$	<i>a</i> = 14.999, <i>b</i> = 0.0353	0.368
Abidin et al.	Clayey silt	25	Power law, w - ρ	<i>a</i> = 121.88, <i>b</i> = -0.363	0.69 - 0.89
2013	SIIt			<i>a</i> = 109.98, <i>b</i> = -0.268	
Osman et al. 2014	Clay	16	Power law, w - ρ	<i>a</i> = 81.12, <i>b</i> = -0.34	0.818
Akinlabi and Adeyemi 2014		7	Linear, <i>PI-p</i>	<i>a</i> = 29.04, <i>b</i> = - 0.002	0.920

 Table 2.2 Examples of empirical studies relating electrical resistivity to the geotechnical properties (between 2012 and 2014)

Notes:

^{*a*} Coefficient of *a*, *b*, and *c* represent constant parameters in the linear (y=a+b.x), power law ($y=a.x^b$), exponential (y=a.exp(b.x)), and quadratic ($y=a.x^2+b.x+c$) regression functions.

Authors	Soil type	No. of data points	Correlation	Parameter values ^a	Coefficient of determination
Lin et al. 2016	Marine		Power law, w - ρ	<i>a</i> = 427.8, <i>b</i> = -1.13	0.930
	clay		Exponential, <i>PI-ρ</i>	<i>a</i> = 124.34, <i>b</i> = -0.239	0.850
			Linear, e -ln(ρ)	<i>a</i> = 4.1663, <i>b</i> = -1.458	0.880
			Quadratic, γ - ρ	<i>a</i> = 0.16, <i>b</i> = -0.0166,	0.720
				<i>c</i> = 20.6	
Jusoh and	Clay		Power law, <i>w-ρ</i>	<i>a</i> = 123.93, <i>b</i> = -0.252	0.816
Osman 2017			Linear, PI -ln(ρ)	<i>a</i> = 29.793, <i>b</i> = -2.71	0.634
Hazreek, et al.	Clayey	25	Power law, w - ρ	<i>a</i> = 110.68, <i>b</i> = -0.347	0.938
2018	silt				
Rezaei et al.		15	Power law, ρ -w	<i>a</i> = 2028.2, <i>b</i> = -1.496	0.68
2018					
Shahandashti et	Clay	842	Linear $\sigma^{0.5}$ -ln(w)	a = -0.3267 $b = 0.215$	0.66
al. 2021	Citay	072	$\operatorname{Linear}, p \operatorname{-in}(w)$	u = -0.5207, 0 = 0.215	0.00

 Table 2.3 Examples of empirical studies relating electrical resistivity to the geotechnical properties (between 2016 and 2021)

Notes:

^{*a*} Coefficient of *a*, *b*, and *c* represent constant parameters in the linear (y=a+b.x), power law ($y=a.x^b$), exponential (y=a.exp(b.x)), and quadratic ($y=a.x^2+b.x+c$) regression functions.

2.3.2. Artificial Intelligence Techniques

Artificial Intelligence (AI) techniques have the potential to revolutionize designs, construction, and maintenance of the infrastructure systems by providing advanced analytics, automation, and predictive capabilities (Zamanian et al. 2023a; Darghiasi et al. 2023a; Baral et al. 2022). In the field of geotechnical engineering, researchers adopted AI techniques such as artificial neural networks and support vector machines to establish relationships between the geotechnical properties and electrical resistivity values. Alsharari et al. (2020) compared the performance of multivariate linear regressions with non-linear regressions and artificial neural networks (ANNs) in quantifying the soil electrical resistivity based on water content, dry unit weight, pore water salinity, and percentage of fine and coarse grains of soils. They found that non-linear regressions perform better than linear regressions in explaining the non-linear and complex interdependencies between the electrical resistivity values and geotechnical properties; however, both models show higher prediction errors than the ANNs. Other researchers also explored the applicability of the ANNs in predicting the soil electrical resistivity based on geotechnical properties. Bian et al. (2015) adopted the ANNs to estimate the electrical resistivity values based on water content, degree of saturation, and porosity. In a similar study, Rashid et al. (2018) performed an experimental study to investigate the variations in the electrical resistivity values of kaolinitedominant clay liners due to variations in water content and dry unit weight. They developed ANNs to predict electrical resistivity values and concluded that ANNs could be used to assess the level of heterogeneity of compacted clay liners.

Samui (2013) examined the application of Support Vector Machines (SVMs) and Least Square Support Vector Machines (LSSVMs) in investigating the associations between electrical resistivity and soil thermal resistivity, coarse-grained fraction, and degree of saturation. He compared the accuracy of the developed SVMs and LSSVMs with ANNs and found that the SVMs and LSSVMs outperform the ANNs, with a slightly better performance of LSSVMs over SVMs. Likewise, Samui (2014) found that the ANNs do not perform as well as Gaussian Process Regression (GPR) in quantifying the soil electrical resistivity values based on the soil thermal resistivity, degree of saturation, and coarse-grained fraction. Although ANNs are more flexible at handling non-linear interactions between the variables, feature extraction and feature engineering are still necessary before training the networks to improve prediction accuracy. In other words, the ANNs cannot derive meaningful features from the unprocessed data due to their shallow structures (Abediniangerabi et al. 2021).

2.4. Gaps in Knowledge

While previous studies have shed light on the correlations between geotechnical properties and electrical resistivity values, there remain significant gaps in understanding the spatial effects of geotechnical properties on electrical resistivity values, as well as the non-linear and complex interactions between them. The following gaps were identified from the literature:

- The presence of spatial autocorrelation between the geotechnical properties and electrical resistivity values in clayey soil has not been studied.
- 2) There is a lack of an analytical tool to extract meaningful information among non-linear and complex relationships between electrical resistivity values and geotechnical properties.

2.5. Research Objectives

The objectives of this research are to:

- Assess the presence of spatial association between electrical resistivity and geotechnical properties. If so, determine the most appropriate spatial regression model to explain the variability of electrical resistivity values considering the spatial effects of geotechnical properties.
- Explore non-linearity and complexity of interactions between the electrical resistivity values and geotechnical properties using artificial intelligence techniques with deep structures such as deep learning model.

The following chapters present the work performed to achieve the research objectives.

CHAPTER 3 EXPLORING SPATIAL ASSOCIATION BETWEEN ELECTRICAL RESISTIVITIES AND GEOTECHNICAL PROPERTIES USING SPATIAL REGRESSION ANALYSIS

Accurate estimation of geotechnical properties and characterization of spatial distributions of geotechnical properties at a site are critical for any successful construction or development activity, especially when considering reliability-based designs such as load and resistance factor design (LRFD) method (Shahandashti et al. 2023). This chapter aims to explore the spatial association between the electrical resistivity values and geotechnical properties (Zamanian and Shahandashti 2023).

3.1. Methodology

3.1.1. Design of Experiments

A full factorial design was established to investigate the effects of water content and dry unit weight on the electrical resistivity of various soil samples with different fine/clay fractions and plasticity indices. A full factorial design generates observations by all possible combinations of factor levels in each complete experiment; it is particularly useful in studying the factor effects when the number of factors is less than five (Zamanian and Yazdandoust 2021 and 2022; Antony et al. 2014; Davim 2012). The water content and dry unit weight were studied at four and three levels. The factors and corresponding factor levels are shown in Table 3.1. The design resulted in at least 12 experimental runs for each soil sample.

Factor	Unit —	Factor Levels			
		1	2	3	4
Water content	%	10	20	30	40
Dry unit weight	kN/m^3	11.8	13.4	14.9	-

Table 3.1 Factors and corresponding factor levels in experimental design

3.1.2. Data Collection

Soil Sample Collection

A total of 44 soil samples were obtained from 13 locations (in four districts) across the state of Texas, US. Figure 3.1 shows the number of boreholes and obtained soil samples on the map of Texas.



Over 50 percent of these areas are underlain by soils with abundant clays of high swelling potential

Less than 50 percent of these areas are underlain by soils with clays of high swelling potential

Less than 50 percent of these areas are underlain by soils with abundant clays of light to moderate swelling potential

These areas are underlain by soils with little to no clays with swelling potential

Figure 3.1 Number of boreholes and collected soil samples on clay map of Texas, US

(Source: Clay map adapted from Olive et al. 1989)

The selected locations are situated in four TxDOT districts in the East, West, South, and North of Texas (Beaumont, Corpus Christi, Fort Worth, and El Paso), representing various TxDOT operational environments and geotechnical conditions. The criteria for selection of these districts include but are not limited to:

- diverse geotechnical characteristics (e.g., soil type, topography, etc.)
- various levels of rainfalls or frequent wetting and drying cycles
- having the most recent projects, which included subsurface investigation (especially those that have problems with the subsurface investigation)

Figures 3.2 to 3.4 illustrates the general soil map, annual average precipitation map, and annual average temperature map of Texas, respectively.



Figure 3.2 General soil map of Texas

(Source: Adapted from Godfrey et al. 1973)


Figure 3.3 Annual average precipitation map of Texas

(Source: Adapted from Spatial Climate Analysis Service, 2000)



Figure 3.4 Annual average temperature map of Texas

(Source: Adapted from Paleontological Research Institution)

Laboratory Testing to Collect Data

Specific Gravity: Specific gravity of soil samples was measured using a water pycnometer according to ASTM D854-14 standard test method. About 50 grams of dried soil material passing the No. 10 (2.00 mm) sieve used in the test. The soil was added to the pycnometer, and the pycnometer was filled about one-half with distilled water. The weights of the empty pycnometer and pycnometer with specimens were measured separately. To remove the entrapped air between the soil particles, a partial vacuum was applied. It is started by applying a low vacuum and then the vacuum level was increased gradually until the water in the flash boils. Then, water was added up to the graduation mark of the pycnometer and weighted. The distilled water was poured in a clean pycnometer, and the combined weight was measured. Using the equations presented in ASTM D54-14, the specific gravity of soil was determined. Figure 3.5 shows the testing procedure on the clayey soil specimens.



Figure 3.5 Specific gravity testing of soil

(Source: Shahandashti et al. 2021)

<u>Atterberg Limits:</u> The performing agency determined the Atterberg limit (liquid limit and plastic limit) of the soil samples according to ASTM D4318-17 standard test method. These tests were conducted on materials passing the No. 40 (0.475-mm) sieve.

Liquid limit is defined as the water content, in percent, of a cohesive soil at the arbitrarily defined boundary between the semi-liquid and plastic states (ASTM D4318-17). First, to conduct the test, small increments of distilled water was added into the soil using a spray bottle to apply a uniform mist of water to the sample. Then, a sufficient amount of soil was placed in the liquid limit device cup, flattened, and finally divided using a grooving tool at the point of maximum thickness. The cup was lifted and dropped at a rate of 2 drops per second until the groove closure was about 13 mm (appropriate water contents should yield to 15 to 35 number of blows). The test

was repeated three times with different water contents. Then to determine the water content, samples were dried in the oven at 100-110 degrees of Centigrade for 24 hours. The water content corresponding to 25 blows was considered as the liquid limit of the soil specimen. Figure 3.6 illustrates the testing procedure using the liquid limit device.



Figure 3.6 Liquid limit testing: (a) the soil is flattened in the device cup, and (b) a groove was made at the center

(Source: Shahandashti et al. 2021)

Plastic limit is defined as the lowest water content, in percent, of a cohesive soil at the boundary between the plastic and semi-solid states (ASTM D4318-17). First, to determine the plastic limit, distilled water was added into the soil and kneaded repeatedly. Then a sufficient

amount of soil was placed on a glass plate and rolled back and forth until threads of about oneeighth inch in diameter (3 mm) were formed and broken into pieces. Then to determine the water content, samples were placed and dried in the oven at 100-110 degrees of Centigrade for 24 hours. The water content corresponding to this stage was considered as the plastic limit of the soil specimen. Figure 3.7 illustrates the rolling device and the state of cracked threats resulted from the experiment.



Figure 3.7 Plastic limit testing (a) Rolling device and (b) cracked and broken threats of 3 mm

(Source: Shahandashti et al. 2021)

<u>Particle Size Distribution:</u> The performing agency determined the particle size distribution of fine-grained soil using the hydrometer method according to ASTM D7928-17 standard test method. The test was performed on material passing the No. 10 (2.0-mm) or finer sieve.

First, approximately 5.0 grams of sodium hexametaphosphate was dissolved in water and added to the sedimentation specimen. The contents were completely mixed with a spatula until all of the soil aggregations are broken-up. The slurry should be soaked overnight (at least 12 hours). Then the slurry was dispersed using a stirring device and transferred into the hydrometer cylinder. A sufficient amount of distilled water was added to bring the level of the water to 1000 ml. Then the cylinder was placed in a constant temperature water bath.

When the soil suspension reaches the temperature of the bath, its contents were completely agitated for about one minute. Then the hydrometer cylinder was placed on the table, and immediately the hydrometer was lowered into the suspension, and the time was recorded. The peak of the meniscus formed on the stem of hydrometer was read to the nearest 0.5 g per liter at the end of two minutes from the time the graduate was set on the table. The cylinder was removed and again placed into the constant temperature bath. The hydrometer readings were obtained at time intervals of 1, 2, 4, 15, 30, 60, 240, and 1440 minutes after the beginning of sedimentation. Figure 3.8 shows the hydrometer test on the clayey soil specimens. Using the equations presented in ASTM D7928-17, particle diameters and the percent finer than a specific diameter were determined.



Figure 3.8 Particle size distribution testing using the hydrometer procedure

(Source: Shahandashti et al. 2021)

Laboratory Electrical Resistivity Test: A four-electrode soil box, current source, resistance measuring equipment, and electrical connections were used to conduct the laboratory testing. First, a specific amount of water was added to the soil and mixed. Then, the soil was placed in the resistivity box and compacted to reach the desired compaction. The soil water contents and dry unit weights were altered from 6 to 45% and 10.2 to 15.7 kN/m3 (60 to 100 pcf), respectively. After the installation of equipment, direct current was applied using two electrodes located at the end of the resistivity box, and the potential drop was measured between two points at the specimen by the AGI SuperSting R8 instrument (ASTM G57-20 2020). The preparation of soil specimens and experimental setup of laboratory resistivity testing are illustrated in Figure 3.9.



Figure 3.9 (a) and (b) preparation of soil specimens, (c) a schematic setup of laboratory electrical resistivity test, and (d) experimental setup of laboratory resistivity test

(Source: Shahandashti et al. 2021)

The measured electrical resistivity of soil is a function of the cross-sectional area of the soil box and electrode spacings (ASTM G57-20 2020) and can be expressed by $\rho = AR/d$, where ρ is electrical resistivity (Ω .m), A is the cross-sectional area of the soil box perpendicular to the current flow (m²), d is the inner distance between the potential electrodes (m), and R is the electrical resistance according to Ohm's law. To eliminate the variability of electrical resistivity measurements because of temperature variations, the measured electrical resistivity values were corrected at a reference temperature of 15.5°C (60°F) using the following equation (ASTM G57-20):

$$\rho_{15.5} = \rho_T \, \frac{(24.5+T)}{40}$$
Eq. 3.1

where $\rho_{15.5}$ is the corrected electrical resistivity at 15.5°C, ρ_T is the electrical resistivity measured at the temperature of *T*°C.

A total of 627 data points were collected from the laboratory physical property tests (e.g., gravimetric water content, Atterberg limits, and specific gravity) and laboratory electrical resistivity tests. Table 3.2 shows the basic statistics for the independent and dependent variables.

Parameters	Abbreviation	Minimum Value	Maximum Value	Mean	VAR	n
Water content	ω (%)	6.6	44.4	23.4	104.2	627
Dry unit weight	$\gamma_d (\mathrm{kN}/m^3)$	10.2	15.7	12.5	1.3	627
Plasticity index	PI (%)	10.6	46.5	28.2	60.7	627
Specific Gravity	G _s	2.6	2.7	2.6	0.0003	627
Electrical resistivity	$ ho (\Omega.m)$	2.3	810.8	24.1	3513.2	627

Table 3.2 Basic statistics for the dependent and independent variables

3.1.3. Spatial Autocorrelation

Spatial autocorrelation (i.e., spatial dependence) is the degree of dependency among similar/dissimilar neighboring observations and mainly emerges when the observations are collected from different locations in space. Linearity, homoscedasticity, independence, and normality are some critical assumptions associated with the linear regression model (Neter et al. 1996). The model assumptions need to be checked before making inferences regarding the model estimates by evaluating the residual plots and performing diagnostic tests such as the Breusch-Pagan test for homoscedasticity, Shapiro-Wilk test for normality, and Moran's I test for spatial autocorrelation. If any of the assumptions are violated, the OLS model is inappropriate and statistical inferences from the model are unreliable (Voss et al. 2006). This study collected data from different locations and investigated the spatial association between the electrical resistivity

values and geotechnical properties. Moran's I test was used to examine the existence of an overall clustering in the OLS regression residuals. The Moran's I test is represented as follows:

$$I = \frac{n}{\sum_{i} \sum_{j} w_{ij}} \frac{\sum_{i} \sum_{j} w_{ij} (x_i - \bar{x}) (x_j - \bar{x})}{\sum (x_i - \bar{x})^2}$$
Eq. 3.2

where *n* is the number of spatial units, *x* is the variable of interest, \bar{x} is the mean of *x*, and w_{ij} is an element of a spatial weight matrix. The spatial weight matrix (**W**) identifies the spatial structure of the observations. Each element of this matrix defines the dependency between two observations (Getis 2009). The spatial weight matrix has different experimental forms based on the geometry of the spatial units, either by their boundaries or distances from each other (Anselin 2005). The selection of a proper weight function is essential to achieve convincing results from spatial modeling, especially when the spatial autocorrelation is strong (Yan-guang 2009; Elhorst 2010). General distance between the locations of the collected soil samples was utilized in this study to identify the neighboring structure of the observations and construct the spatial weights. Figure 3.10 shows an example of the data points arrangement (i.e., borehole locations) and neighboring structure of the observations based on their distances.



Figure 3.10 An example of the data points arrangement (i.e., borehole locations) and neighboring structure of the observations based on their distances.

The distance-based weight matrices are the most appropriate form for a data set with point locations (Anselin and Sergio 2014). If " d_{ij} " denotes the distance between the location of *i* and *j*, and "*d*" indicates a distance threshold where direct spatial influence between the observations no longer exists, the spatial weights of the corresponding weight matrix are constructed as follows (Chen 2012):

$$w_{ij} = \begin{cases} 1 & , & 0 \le d_{ij} \le d \\ 0 & , & d_{ij} > d \end{cases}$$
 Eq. 3.3

which gives a binary matrix of 0 and 1. Typically, there is no unique approach to determine the threshold distance for identifying the neighboring locations (Walker et al. 2000; Anselin 2005). The most widely used approaches are to assess the robustness of estimated spatial regression models and the magnitude of Moran's I for a series of threshold distances. The distance at which the model shows the maximum log-likelihood value, highest Moran's I value, highest pseudo-Rsquared, and lowest residual standard error is determined as the appropriate threshold distance (Wang et al. 2007; Chi and Zhu 2008; Stakhovych and Bijmolt 2009; Elhorst 2010). The other approach is to identify the threshold distance by creating a semi-variogram of the variables (Hession and Moore 2011). The off-diagonal elements of the weight matrix with non-zero values denote the dependency of the neighboring observations. However, the diagonal elements of the weight matrix represent the self-influence of the observations that were excluded from the spatially lagged variables (i.e., diagonal elements of the weight matrix were set to zero). Then the weight matrix was standardized using a row-normalization approach in which all the weights in each row sum to unity ($\sum_{j=1}^{n} w_{ij} = 1$).

The null hypothesis of the Moran's I test is that the regression residuals are randomly distributed in space. By rejecting the null hypothesis, it is concluded that there is evidence of spatially autocorrelated residuals (alternative hypothesis). Ignoring the presence of spatial dependence in the OLS model leads to underestimation or overestimation of actual variance in the case of positive and negative dependence, respectively, which consequently affects the significance of the model (Schabenberger and Gotway 2005; Cressie 2015). Moran's I value ranges from -1 and +1, and its significance is evaluated using a P-value and z-score. The negative values represent the clustering between dissimilar values, while positive values represent the clustering between dissimilar values, while positive values represent the autocorrelation in the regression residuals, and the residuals are randomly distributed.

3.1.4. Spatial Regression Analysis

The spatial dependence between the observations is accounted into a regression model using the spatial weight matrix by three methods; (1) inclusion of the effect of a change in the dependent variable of one location on the dependent variable of a neighboring location (endogenous interaction effects), (2) inclusion of the effect of a change in the independent variables of one location on the dependent variable of a neighboring location (exogenous interaction effect), and (3) inclusion of the effect of dependency in the residuals in one location on a neighboring location (Calderon 2009). In this study, three spatial regression models were examined: Spatial Durbin Model (SDM), Spatial Lag or Autoregressive Model (SAR), and Spatial Error Model (SEM). The SDM is a general model that includes both exogenous and endogenous interaction effects and has the form of:

$$\rho_{15,5} = \eta W \rho_{15,5} + X \beta + W X \theta + \varepsilon$$
 Eq. 3.4

where $\rho_{15.5}$ is an $(n \times 1)$ vector of observations on the corrected electrical resistivity at 15.5°C (dependent variable), X is an $(n \times k)$ matrix of observations on the geotechnical engineering parameters (independent variables), W is an $(n \times n)$ matrix of spatial weight, β is a $(k \times 1)$ vector of regression parameters, η is a coefficient on the spatially lagged dependent variable, θ is a $(k \times 1)$ vector of the spatially lagged independent variable, and ε is an $(n \times 1)$ vector of independently and identically normally distributed errors. In this research, to avoid multicollinearity in the analyses, the geotechnical parameters (degree of saturation, liquid limit, and void ratio) with the lowest significant test statistics that have a high correlation with the other variables were removed from

the model. The simple linear regression used gravimetric water content, and the multiple linear regression used gravimetric water content, dry unit weight, and plasticity index as independent variables to explain the variability of electrical resistivity in the analyses. The SAR model only includes the endogenous interaction effects (θ =0 in equation 4) and is expressed as:

$$\rho_{15.5} = \eta W \rho_{15.5} + X \beta + \varepsilon$$
 Eq. 3.5

where the variables are defined as the same for the SDM model. In the SDM model, the beta coefficients of the SAR model are not represented by partial derivatives along the diagonal (Golgher and Voss 2016). Therefore, the changes in the dependent variable in one location due to a one unit increase in the independent variables in the same location (direct effect) and another location (indirect effect) are calculated for SAR models. The direct, indirect, and total effects (i.e., direct and indirect effects) can be presented by the following equations:

Direct Effect =
$$\frac{1}{n} tr(S_k(W))$$
 Eq. 3.6

$$Total \ Effect = \frac{1}{n} \ i'_n (S_k(W)) i_n$$
 Eq. 3.8

where $tr(S_k(W))$ is the trace of the partial derivative matrix for variable *k*, *n* is the number of spatial units, and *i_n* is the identity matrix. In contrast, the spatial dependence in the SEM is modeled only by the spatially lagged error terms and considers neither the exogenous nor endogenous interaction effects (θ =0 and η =0 in equation 4), which has the form of:

$$\rho_{15.5} = X\beta + u$$
, $u = \lambda Wu + \varepsilon$ Eq. 3.9

where u is an $(n \times 1)$ vector of error terms and λ is a spatial error lag coefficient. The Lagrange Multiplier (LM) tests were performed on the OLS residuals to decide whether the spatial lag (SAR) or spatial error model (SEM) is the most appropriate model for the analysis of the data (Anselin 2005). There are four LM test statistics: standard LM-Error, standard LM-Lag, Robust LM-Error, and Robust LM-Lag. First, the standard LM tests are performed, and then the model with the significance test statistic is selected. If neither of the tests is significant, it indicates that the OLS model is more appropriate. However, if both standard LM tests are significant, which commonly happens in practice, the Robust forms of LM test are used, and the model with the (most) significance test statistic is selected as the most appropriate model (Anselin 2005). Another approach is to start with the widely used model (i.e., SDM) if there is a global effect (Lesage 2014). Then to further evaluate the goodness-of-fit of the nested models (when a complex model can be reduced to a simpler model by restricting certain parameters), the Likelihood Ratio (LR) test was utilized (Anselin 2005). The null hypothesis of the test is that a complex model should be reduced to a simpler model by restricting some of the model parameters. By rejecting the null hypothesis, it is concluded that the complex model is more appropriate and should not be restricted to the

simpler model (alternative hypothesis). Log-Likelihood (LIK), Bayesian Information Criterion (BIC) (Schwarz Information Criterion), and Akaike's Information Criterion (AIC) were also used to compare the performance of the non-nested models (Yang and Fik 2014; LeSage 2014). The model with the highest LIK and lowest AIC or BIC was considered the best model that fits the data.

3.2. Results

3.2.1. Results of Standard Regression Analysis

The simple and multiple linear regression models were fitted to the electrical resistivity data to test the performance of OLS models in defining a relationship between the geotechnical parameters and electrical resistivity values and checking the model assumptions. The simple and multiple linear regression models developed using the original data are as follows:

$$\rho = 77.80 - 2.293 \,\omega$$
 Eq. 3.10

$$\rho = 283.79 - 2.742 \,\omega - 16.840 \,\gamma_d + 0.533 \,PI$$
 Eq. 3.11

where ρ is electrical resistivity, ω is water content, γ_d is dry unit weight, and *PI* is plasticity index. Table 3.3 presents the results of the fitted standard regression models (OLS) before and after transforming the electrical resistivity values.

Table 3.3 Summary of results of OLS model before and after transforming the electrical

	Multiple Linear Regression		Simple Linea	ar Regression
	OLS with no transformation	OLS using Box-Cox transformation	OLS with no transformation	OLS using Box-Cox transformation
Intercept	283.79**	-0.524**	77.80**	0.093**
Water Content	-2.742**	0.011**	-2.293**	0.010**
Dry Unit Weight	-16.840**	0.038**		
Plasticity Index	0.533*	0.004**		
Adjusted R- squared	0.26	0.76	0.15	0.62
Standard Error of Residual	51.05	0.067	54.5	0.084
LIK	-3353.54	807.23	-3395.57	661.37
AIC	6717.08	-1604.47	6797.14	-1316.73
BIC	6739.29	-1582.27	6810.46	-1303.41
No. of Observations	627	627	627	627

resistivity values

Notes: '*' indicates the significance at the 5% level and '**' indicates the significance at the 1% level. "OLS" denotes Ordinary Least Squares, "LIK" denotes Log-Likelihood, "AIC" denotes Akaike's Information Criterion, and "BIC" denotes Bayesian Information Criterion.

According to Table 3.3, from the initial analysis using multiple linear regression with no transformation, the regression coefficients show statistically significant negative values for gravimetric water content and dry unit weight (at the 5% significance level) and a positive value for plasticity index (at the 1% significance level). The negative regression coefficients imply that the gravimetric water content and dry unit weight have inverse relationships with the electrical resistivity values. For example, a unit increase in the gravimetric water content results in a 2.742 decrease in the electrical resistivity, keeping other independent variables constant. However, the positive regression coefficient for the plasticity index shows a direct relationship with the electrical resistivity value, which is inconsistent with the literature. Figures 3.11 and 3.12 illustrate the residual plots of the OLS regression model with no transformation of the electrical resistivity values. The presence of a funnel in the plot of residuals versus fitted values and skewness in the normal probability plot are indications of heteroskedasticity and non-normality of the error terms, respectively. Besides, as shown in Table 3.4, the Breusch-Pagan test and Shapiro-Wilk test show that the assumptions of homoskedasticity and normality of the linear regression model are not satisfied (rejection of null hypotheses at the 10% level of significance). The results of simple linear regression agree with those of multiple linear regression. Therefore, the electrical resistivity values were transformed using the Box-Cox transformation to stabilize the error variance and mitigate the problem of non-normality of the error terms.



Figure 3.11 Residuals versus fitted values for the OLS model with no transformation on the electrical resistivity values.



Figure 3.12 Normal probability plot for the OLS model with no transformation on the electrical

resistivity values.

 Table 3.4 Diagnostic test results for multiple linear regression before transformation

Test	Value	Prob
Shapiro-Wilk	0.4710	0.0000
Breusch-Pagan	29.311	0.0000

The regression analysis using a Box-Cox transformation yielded the following equations:

$$\rho^{-0.5} = 0.093 + 0.010 \,\omega \qquad \qquad \text{Eq. 3.12}$$

$$\rho^{-0.5} = -0.524 + 0.011 \,\omega + 0.038 \,\gamma_d + 0.004 \,PI$$
 Eq. 3.13

The results of the OLS analysis with the transformed dependent variable (λ =-0.5) are presented in Table 3.3. The regression coefficients for gravimetric water content, dry unit weight, and plasticity index show statistically significant positive values (at the 5% level of significance). Note that the positive signs mean that the independent variables directly correlate with the inversed electrical resistivity values. In other words, the results imply that the gravimetric water content, dry unit weight, and plasticity index have significant inverse relationships with the electrical resistivity value, which is consistent with the literature. Figures 3.13 and 3.14 illustrate the residual plots of the OLS regression model using Box-Cox transformation on the electrical resistivity values. No clear pattern can be observed in the plot of residuals versus fitted values shown in Figure 3.13. Table 3.5 presents the diagnostic test results for multiple linear regression after transformation. The Breusch-Pagan test also shows that the assumption of the constant variance of error terms is satisfied after transformation at the 1% level of significance. The skewness of data in the normal probability plot is removed after transforming the electrical resistivity values; however, it can be observed that the residuals are less spread than the normal distribution (lighter-tailed). The Shapiro-Wilk test also confirms that the normality assumption is violated (rejection of null hypothesis at the 10% level of significance). Although the non-normality of error terms has remained even after transformation, no more transformations were used since the OLS model is relatively robust to non-normality in the absence of skewness (Neter et al. 1996). The diagnostic tests showed similar results for the simple linear regressions.



Figure 3.13 Residuals versus fitted values for the OLS model with transformed electrical

0 0000 2 Standardized Residuals 0 1 2 n -3 -2 -1 0 1 2 3 Theoretical Quantiles

resistivity values.

Figure 3.14 Normal probability plot for the OLS model with transformed electrical resistivity

values.

Test	Value	Prob
Shapiro-Wilk	0.9857	0.0000
Breusch-Pagan	27.185	0.0101*

 Table 3.5 Diagnostic test results for multiple linear regression after transformation

Note: '' indicates the significance at the 1% level.*

The Moran's I test provides strong evidence of positive spatial autocorrelation (Table 3.6) in the regression residuals, implying that the residuals are not independently distributed. According to Table 3.6, although the Moran's I value for multiple linear regression is greater than its value for simple linear regression, the higher z-score for simple linear regression indicates a stronger spatial autocorrelation between the regression residuals for simple than multiple linear regression (i.e., the greater the z-score, the stronger the spatial autocorrelation). Since the assumption of independence of the linear regression is violated, the OLS model might be an inappropriate approach to quantify the relationship between electrical resistivity and geotechnical properties. Besides, any statistical inferences regarding the coefficient estimates might be unreliable. Therefore, the spatial regression models were examined to account for the spatially autocorrelated residuals.

	Multiple Linear Regression	Simple Linear Regression
Moran's I	0.281	0.152
P-value	0.000*	0.000*
z-score	32.872	44.714

 Table 3.6 Summary of results of Moran's I tests for the OLS residuals

Notes: '*' indicates the significance at the 1% level.

3.2.2. Results of Spatial Regression Analysis

The simple and robust forms of Lagrange Multiplier tests (LM) were used on the OLS results to determine the most appropriate spatial model for the analysis. Table 3.7 presents a summary of the results of LM tests for the simple and multiple linear regression models. Since both simple tests (LM error and LM lag) are highly significant and suggest using the spatial regression models, the robust form of LM error and LM lag tests were tested. The robust tests also show highly significant values for both SAR and SEM; however, it appears that the test statistic for the spatial error model (SEM) is more significant.

	Result			
Test	Multiple Linear Regression	Simple Linear Regression		
LM error	935.47*	1363.10*		
LM lag	306.75*	814.26*		
Robust LM error	629.29*	557.77*		
Robust LM lag	30.567*	8.90*		

Table 3.7 Summary of results of Lagrange Multiplier tests for the OLS residuals

Notes: '*' indicates the significance at the 1% level.

Although it is concluded from the LM tests that the SEM is the most appropriate model, the SAR and SDM were also examined to compare the performance of different spatial regression models. The results of these spatial analyses with the transformed electrical resistivity values are summarized in Table 3.8.

	Multiple Linear Regression		Simple	Simple Linear Regression		
	SDM	SAR	SEM	SDM	SAR	SEM
Intercept	-0.020	-0.731*	-0.534*	-0.021	-0.159*	0.094*
Water Content	0.011*	0.011*	0.011*	0.010*	0.010*	0.010*
Dry Unit Weight	0.042*	0.041*	0.042*			
Plasticity Index	0.003*	0.004*	0.003*			
Lag. Water Content	-0.009*			-0.007		
Lag. Dry Unit Weight	-0.045*					
Lag. Plasticity Index	-0.002*					
$\boldsymbol{\eta}$ / λ Coefficient	0.765	0.576	0.817	0.800	0.749	0.808
Pseudo-R-squared	0.83	0.82	0.83	0.67	0.67	0.67
Standard Error of Residual	0.055	0.040	0.044	0.082	0.070	0.078
LIK	906.91	888.87	904.45	704.97	703.79	704.88
AIC	-1795.8	-1765.7	-1796.9	-1399.95	-1399.58	-1401.76
BIC	-1755.8	-1739.1	-1770.3	-1377.74	-1381.81	-1384.00
LM Test for Residual Autocorrelation	5.25*	20.85	1.54**	114.68	3.30*	6.96*
No. of Observations	627	627	627	627	627	627

Table 3.8 Summary of results of spatial regression models with the transformed data

Notes: '*' indicates the significance at the 1% level, '**' indicates the significance at the 10% level. "SDM" denotes Spatial Durbin Model, "SAR" denotes Spatial Lag Model, "SEM" denotes Spatial Error Model, " η " denotes coefficient of spatially lagged dependent variable, " λ " denotes coefficient of spatial error lag, "LIK" denotes Log-Likelihood, "AIC" denotes Akaike's Information Criterion, and "BIC" denotes Bayesian Information Criterion.

The SDM which includes both exogenous and endogenous interaction effects is presented by the following equations:

$$\rho_i = -0.020 + 0.765 \, W\rho + (0.011 - 0.009) \, \omega + (0.042 - 0.045 \, W) \, \gamma_d + (0.003 - 0.002 \, W) \, PI$$
Eq. 3.14

where *W* is the average value in neighboring locations. According to Table 3.8, for multiple linear regression, the spatial lag coefficient of the SDM is positive, meaning that a change in the electrical resistivity of one location has positive effects on the electrical resistivity values of neighboring locations. These effects decay as moving toward higher-order neighbors. In other words, the variations of electrical resistivity values in one location influence the electrical resistivity of nearby locations more than further locations. The likelihood ratio test and Wald statistics show that the spatial lag coefficient of the SDM (η =0.765) is significant at the 1% level.

Similarly, the SAR models were developed and can be estimated as follows:

$$\rho_i = -0.731 + 0.576 W \rho + 0.011 \omega + 0.041 \gamma_d + 0.004 PI$$
 Eq. 3.15

For multiple linear regression, the spatial lag coefficient of the SAR model is positive and significantly different from zero at the 1% level (η =0.576). The SAR model presents positive but lower spillover effects in the neighboring locations rather than the SDM.

The SEMs are presented by the following equations:

$$\rho_i = -0.534 + 0.011 \,\omega + 0.042 \,\gamma_d + 0.003 \,PI + 0.817 \,Wu$$
 Eq. 3.16

where Wu is the average error of prediction in neighboring locations. The spatial error lag coefficient of the SEM is positive and significantly different from zero at the 1% level (λ =0.809). The spatial error lag coefficient of the SEM shows the strength of spatial autocorrelation among the error terms meaning that the unexplained variabilities of the electrical resistivity values follow a systematic distribution in space. The results of the simple linear regression show similar patterns to multiple linear regression results.

The signs and magnitudes of the SEM model parameters are similar to the standard regression models. The SEM model parameters are also highly significant for the three geotechnical parameters (gravimetric water content, dry unit weight, and plasticity index). Since the coefficients of the SAR model do not accurately explain the effects of geotechnical properties on electrical resistivity, a direct comparison of the regression parameters of the SAR model and standard regression model is inappropriate (LeSage and Dominguez 2012). Therefore, the average direct, indirect, and total effects of a change in each of the three geotechnical parameters on the electrical resistivity were calculated for the SAR model and summarized in Table 3.9.

Variable	Direct effect	Indirect effect	Total effect	P-value
Water Content	0.011	0.015	0.026	0.000*
Dry Unit Weight	0.041	0.055	0.096	0.000*
Plasticity Index	0.004	0.005	0.009	0.000*

 Table 3.9 Average effects of explanatory variables on the electrical resistivity values for the

 SAR model

Note: '' indicates the significance at the 1% level.*

According to Table 3.9, the corresponding direct effects of gravimetric water content, dry unit weight, and plasticity index are smaller than the indirect effects, holding the same signs, which are associated with the transformation used on the electrical resistivity values. Similar to the standard regression model, the total effects of geotechnical properties on the electrical resistivity value are positive and highly significant at the 1% level. It again implies that an increase in the geotechnical properties has a decreasing effect on the electrical resistivity values. A noticeable difference is that the coefficients of the geotechnical parameters in the SAR model are shifted toward positive values compared to the standard regression model due to considering both direct and indirect effects. The coefficient variations imply that the variability of electrical resistivity is less influenced by the variation of geotechnical properties while considering the spatial effects.

The results of Lagrange Multiplier diagnostic tests for the spatial dependence of multiple linear regression show that the SEM and SDM models removed the problem of spatially autocorrelated residuals at the 10% and 1% level of significance. However, the spatial autocorrelation has remained in the SAR residuals (the null hypothesis is rejected at the 1% level of significance). Moreover, the likelihood ratio (LR) test was utilized to evaluate the goodness-offit of the nested models (i.e., SAR and OLS, or SEM and OLS). The test results show that the SAR and SEM models outperform the standard regression models at the 1% level of significance, and they should not be restricted to a simpler model (i.e., the OLS model). Comparing the LIK, AIC, and BIC statistics, it appears that the SEM is a better fit for the electrical resistivity data compared to the SDM. Therefore, according to the diagnostic tests and statistics, it is concluded that the spatial error model (SEM) is the best spatial model compared to the SDM and SAR models. Besides, the SEM provides more accurate estimates of the regression parameters in comparison to the standard regression model due to considering the spatial effects in the analysis.

3.2.3. Robustness of Spatial Regression Models Based on Threshold Distance

In this paper, threshold distances of 0.4 km (0.25 mi), 0.8 km (0.5 mi), 1.2 km (1 mi), 1.6 km (2 mi), 6.4 km (4 mi), 9.6 km (6 mi), 12.8 km (8 mi), 16.1 km (10 mi), 32.2 km (25 mi), 48.3 km (30 mi), 80.5 km (50 mi), and 160.9 km (100 mi) were examined to construct spatial weight matrices to assess the robustness of spatial regression models and investigate the spatial autocorrelation in the regression residuals of electrical resistivity data (at shorter threshold distances than 0.4 km (0.25 mi), no neighbor was found for some locations). Table 3.10 and 3.11 represent the values of log-likelihood, pseudo-R-squared, residual standard error, and Moran's I of OLS residual considering different threshold distances for the SEM and SAR, respectively. Although the log-likelihood of the SEM shows more variation than the SAR model, its value decreases as the threshold distance increases in both models. The log-likelihood has the highest

value at 0.4 km (0.25 mi) threshold distance in both models. Similarly, the value of Moran's I decreases as the threshold distance increases and has the highest value at 0.4 km (0.25 mi) threshold distance. The pseudo-R-squared and residual standard error have approximately constant values at different lag distances. Therefore, a threshold distance of 0.4 km (0.25 mi) was determined to construct the spatial weights and perform the spatial regression analyses on the electrical resistivity data based on the highest log-likelihood, highest Moran's I, highest pseudo-R-squared, and lowest residual standard error.

Spatial	Threshold Distance in km	Log-	Pseudo-R-	Residual	Moran's I
Model	(mi) ^a	Likelihood	Squared	Standard Error	OLS Residual
	0.4 (0.25)	904.45	0.83	0.0437	0.280*
	0.8 (0.5)	882.38	0.81	0.0586	0.241*
	1.2 (1)	889.73	0.82	0.0512	0.236*
	1.6 (2)	884.19	0.81	0.0526	0.212*
SEM	6.4 (4)	884.19	0.81	0.0526	0.218*
	9.6 (6)	884.12	0.81	0.0526	0.212*
	12.8 (8)	884.89	0.81	0.0529	0.211*
	16.1 (10)	884.89	0.81	0.0529	0.211*
	32.2 (25)	884.89	0.81	0.0529	0.211*
	48.3 (30)	884.89	0.81	0.0529	0.210*
	80.5 (50)	884.89	0.81	0.0529	0.210*
	160.9 (100)	873.37	0.80	0.0612	0.194*

Table 3.10 Variations of log-likelihood, pseudo-R-squared, residual standard error, and Moran'sI of OLS residual considering different threshold distances for the SEM

Notes: '*' indicates the significance at the 1% level. "OLS" denotes Ordinary Least Squares, "SAR" denotes Spatial Lag Model, and "SEM" denotes Spatial Error Model.

^{*a*} numbers in parentheses represent threshold distances in miles.

Spatial	Threshold Distance in km	Log-	Pseudo-R-	Residual	Moran's I
Model	(mi) ^a	Likelihood	Squared	Standard Error	OLS Residual
	0.4 (0.25)	888.87	0.82	0.0401	0.280*
	0.8 (0.5)	879.00	0.81	0.0457	0.241*
	1.2 (1)	886.29	0.81	0.0437	0.236*
	1.6 (2)	881.65	0.81	0.0451	0.212*
SAR	6.4 (4)	881.65	0.81	0.0451	0.218*
	9.6 (6)	881.65	0.81	0.0451	0.212*
	12.8 (8)	883.67	0.81	0.0493	0.211*
	16.1 (10)	883.67	0.81	0.0493	0.211*
	32.2 (25)	883.67	0.81	0.0493	0.211*
	48.3 (30)	883.67	0.81	0.0493	0.210*
	80.5 (50)	883.67	0.81	0.0493	0.210*
	160.9 (100)	875.66	0.80	0.0679	0.194*

Table 3.11 Variations of log-likelihood, pseudo-R-squared, residual standard error, and Moran'sI of OLS residual considering different threshold distances for the SAR

Notes: '*' indicates the significance at the 1% level. "OLS" denotes Ordinary Least Squares, "SAR" denotes Spatial Lag Model, and "SEM" denotes Spatial Error Model.

^{*a*} numbers in parentheses represent threshold distances in miles.

CHAPTER 4 INVESTIGATING COMPLEX RELATIONSHIP BETWEEN ELECTRICAL RESISTIVITY VALUES AND GEOTECHNICAL PROPERTIES USING DEEP LEARNING

Linear regression analysis and artificial intelligence techniques with shallow structures cannot discover non-linear and complex relationships between electrical resistivity values and geotechnical properties. Deep learning models can capture highly non-linear relationships in large datasets because of their algorithm's flexibility (Jalal et al. 2021). However, the deep learning models' generalization capabilities are in doubt due to overparameterization in which a trained deep learning model can overfit and assign overconfident predictions (Maronas et al. 2021). This chapter proposes an approach to extracting meaningful information from the non-linear and complex relationship between electrical resistivity and geotechnical properties (Zamanian et al. 2023c).

4.1. Methodology

The proposed approach includes two main steps: (1) collecting data from laboratory soil physical property and electrical resistivity tests and (2) training a deep learning model using the obtained laboratory data.
4.1.1. Data Collection

To introduce more variability in the previous dataset used in the spatial regression analysis, an additional 19 soil samples were collected and tested to identify the influencing geotechnical properties on the electrical resistivity values in clayey soil. Figure 4.1 shows the locations of soil sample collection on the Texas map. These sites are located within a distance of up to 420 miles.



Figure 4.1 Locations of the soil sample collection on the Texas map

(Source: Adapted from Olive et al. 1989)

Based on the experimental design, each soil sample was mixed with different amounts of water and compacted in a soil box with three compaction efforts to conduct the electrical resistivity tests. A total of 842 laboratory measurements were conducted using the AGI SuperSting R8 instrument following the standard test method for measuring electrical resistivity using the Wenner four-electrode method (ASTM G57-20 2020). In addition to the laboratory electrical resistivity measurements, soil physical property tests were conducted to quantify the plasticity index, fine fraction, clay fraction, and specific gravity of the soil samples. Table 4.1 summarizes the basic statistics (e.g., range of values, mean, variance) of the measured parameters. According to Table 4.1, the soil samples are classified as low (CL) to high (CH) plasticity clayey soils based on the unified soil classification system (USCS).

Parameters	Abbreviation	Minimum Value	Maximum Value	Mean	VAR	n
Water content	ω (%)	6.6	66.0	24.1	125.7	842
Dry unit weight	$\gamma_d (\mathrm{kN}/m^3)$	7.8	15.8	12.7	1.3	842
Plasticity index	PI (%)	6.5	46.5	28.4	65.6	842
Fine fraction	F (%)	54.2	98.0	85.7	94.5	842
Clay fraction	C (%)	9.4	68.8	43.3	188.2	842
Specific Gravity	G _s	2.6	2.72	2.65	0.001	842
Electrical resistivity	$\rho\left(\Omega.m ight)$	2.3	995.0	27.3	5394.1	842

Table 4.1 Basic statistics of the input and output parameters

4.1.2. Deep Learning Model

Artificial neural networks (ANNs) mimic the biological learning mechanism of the human brain, enabling the exploration of intricate and non-linear associations between input and output features (Abediniangerabi et al. 2021). The neural networks are composed of one input and output layer, and one or multiple hidden layers. Each layer comprises one or more neurons, and each layer's neurons are interconnected to neurons at the next layer by weighted connections. The neurons process elements of a neural network and resemble human brain cells (Darghiasi et al. 2024). The networks that possess multiple hidden layers are referred to as "Deep Learning" or "Deep Neural Networks" (DNNs). Figure 4.2 shows the structure of a deep learning model with four input features (geotechnical properties) and one output feature (electrical resistivity). The deep learning models use simple but non-linear algorithms to extract multiple higher levels of representation from the raw data and reveal complex patterns (LeCun et al. 2015). While shallow and deep neural networks possess the universal approximation property, research shows that deep network architectures (i.e., networks with two or more hidden layers) perform better than shallow network architectures (i.e., networks with one hidden layer) with an exponentially lower number of training parameters (Bengio and LeCun, 2007).



Notes: Input layer: " ω " denotes water content, " γ_d " denotes dry unit weight, "PI" denotes plasticity index, and "F" denotes fine fraction. Output layer: " ρ " denotes electrical resistivity

Figure 4.2 Structure of a deep learning model with four input features and an output feature

The neural networks represent and compute the non-linear associations between the input and output features in the hidden layers (Erzin et al. 2010). Each hidden layers' neuron uses a nonlinear activation function to establish a relationship between the input and output features. Mathematically, a combination of the non-linear weighted sum is approximated by Equation 4.1.

$$f(\mathbf{X}) = \phi(W^T \mathbf{X} + b)$$
Eq. 4.1

where X is the matrix of input features, ϕ is the activation function, W is the vector of neuron's weights, and b is the bias term at a hidden layer. The commonly used activation functions in hidden layer's neurons include rectifier linear unit (ReLU), hyperbolic tangent, and sigmoid functions. Computationally, the ReLU function exhibits a faster learning rate than sigmoid and hyperbolic tangent functions in deep neural networks (Su et al. 2017) since it does not saturate the output to a given value by increasing or decreasing the input features (Achieng 2019). This study will employ the ReLU function in all hidden layers except for the last hidden layer. Instead, a linear function was used to connect the last hidden layer to the output layer.

A backpropagation learning algorithm is extensively exploited for training the networks, involving a feedforward and a backward process (Erzin et al. 2010). In the feedforward process, the past observations are fed into the input layer and then propagated to the hidden layers to extract further information. The connection weights are determined to transfer the extracted information to the output layer to predict the outputs. The backward process updates the connection weights and biases based on the predicted outputs and actual values. The training cycle of networks is repeated many times to find the optimal set of weights that will yield the optimal output for any input (Caglar and Arman 2007).

In this study, deep learning models will be trained by four input geotechnical properties including water content, dry unit weight, plasticity index, and fine fraction to estimate the output that is the soil electrical resistivity values. The data was standardized by subtracting the mean from all observations and then scaling to unit variance to speed up the model training, which is essential when dealing with a large volume of data (Jayalakshmi and Santhakumaran 2011). Then the data were randomly split into two sets by a ratio of 80 to 20 before model training. In other words, 80 percent of the observations were used to train the deep learning models, and the remaining 20

percent of the observations were used to evaluate the developed model accuracies. The number of hidden layers for deep learning model shall be chosen appropriately to ensure the best accuracy of the model (Choldun et al. 2019). The maximum number of neurons in the hidden layers is determined by 2I+1, where I represents the number of input variables (Caudill 1988). Therefore, the optimum number of neurons for the hidden layers was selected based on a trial-and-error by altering the number of neurons from one to nine. Note that selection of many hidden layers and hidden layer's neurons may lead to overfitting (i.e., high variance) if the level of complexity of the problem is disregarded (Uzair and Jamil 2020). The optimal model was then selected based on the minimum root mean square of errors (RMSE) and mean absolute errors (MAE) for the testing dataset using 100 iterations. For model's hyperparameter tuning (to mitigate any potential overfitting and assess the model performance to new unseen data), cross-validation was used during model training (Hastie et al. 2009). An n-fold cross-validation method divides the training data into *n* equal parts (i.e., folds). The model is then trained based on *n*-1 parts and then evaluated on the remaining part. The process is repeated *n* times until every part was used once for validation. Then the final model performance is calculated by averaging the results of each iteration. The hyperparameters of the model, such as epoch number and batch size, was chosen based on a grid search using 10-fold cross-validation with a learning rate of 0.001. Table 4.2 presents the ranges of the model hyperparameters that will be used in the grid search. The batch size was also examined in conjunction with the execution time of the training process (Abediniangerabi et al. 2021).

Hyperparameters	Values
Number of neurons in different layers	1 to 9
Number of epochs	10, 20, 30, 40, 50, 100, 150, 200, 300, 400, 500
Batch sizes	64, 128
Activation function	ReLU, Linear
Optimizer	Adam
Learning rate	0.001
Losses	RMSE, MAE

Table 4.2 Model hyperparameters' ranges in grid search

As a result, a deep learning model will be adopted to evaluate the applicability of DNNs in determining the associations between the geotechnical properties and electrical resistivity values in clayey soils. The performance of the trained deep learning model shall be compared to artificial neural networks, support vector machines, and multivariate linear regressions.

4.2. Results

4.2.1. Descriptive Statistics

This section provides descriptive analyses of the data obtained by laboratory experiments to identify the most influencing factors affecting electrical resistivities. Figures 4.3 to 4.8 illustrate the frequency histograms of the measured geotechnical properties for the soil samples. Accordingly, all geotechnical properties show a wide range of values (i.e., high variance), except for the specific gravity. The high variance in the geotechnical properties (i.e., water content, plasticity index, dry unit weight, fine fraction, and clay fraction) indicates that these variables can be useful in determining the variance of the electrical resistivity.



Figure 4.3 Water content frequency distribution



Figure 4.4 Dry unit weight frequency distribution



Figure 4.5 Plasticity index frequency distribution



Figure 4.6 Clay fraction frequency distribution



Figure 4.7 Fine fraction frequency distribution



Figure 4.8 Specific gravity frequency distribution

Moreover, Spearman's correlation analysis was performed to describe the strength of the pairwise relationships between electrical resistivities and geotechnical properties. Spearman's correlation tests for a monotonic dependence between the ranked values of two variables, without assuming linearity of the relationship (Deebani and Kachouie 2022). Figure 4.9 presents Spearman's correlation coefficients for the electrical resistivity and geotechnical properties on a heatmap. Spearman's coefficients range between -1 and 1. A positive coefficient indicates monotonic changes in the same direction, whereas a negative coefficient shows monotonic changes in the opposite direction (Schober et al. 2018). According to Figure 4.9, electrical resistivity shows a strong negative correlation with the water content ($r_s = -0.77$). In other words, the electrical resistivity of soil significantly decreases with an increase in the water content. The literature also confirms that the water content inversely influences the electrical resistivity of clayey or sandy soils (Zamanian and Shahandashti 2022; Alsharari et al. 2020). Based on Figure 4.9, the electrical resistivity shows weak correlations with dry unit weight, plasticity index, fine fraction, and clay

fraction. Note that Spearman's correlation solely measures the degree of a monotonic relationship between two variables. However, there might be strong non-monotonic relationships between the variables which cannot be captured by Spearman's correlation (Griessenberger et al. 2022).



Figure 4.9 Spearman's correlation coefficient heatmap of the electrical resistivity and

geotechnical properties

4.2.2. Deep Learning Results

A deep learning model consisting of three hidden layers was constructed, with each layer containing 8, 8, and 9 neurons, respectively. Four geotechnical properties such as water content, plasticity index, dry unit weight, and fine fraction were fed into the neurons of the input layer to estimate the soil electrical resistivities. Figures 4.10 and 4.11 illustrate the training and testing loss functions (mean squared and absolute errors) for the developed deep learning model across different epochs, ranging from 1 up to 500. The epoch upper limit was set to the highest possible value to ensure that the loss functions converge during the training process (Abediniangerabi et al. 2021). According to Figures 4.10 and 4.11, the deep learning loss functions converge to a constant value as the number of epochs increases. The MSE and MAE loss functions start to saturate around the same value and converges in about 100 epochs. The fluctuations observed on the training and testing and testing curves do not affect the model's overall accuracy (Zhang et al. 2021).



Figure 4.10 MSE loss function for the training and testing datasets



Deep Learning Model Loss Function (MAE)

Figure 4.11 MAE loss function for the training and testing datasets

Validating black-box models such as deep learning requires an understanding of the underlying relationships between input and output variables, which can be achieved by extracting features' importance (Zhong et al. 2021). Figure 4.12 illustrates the relative importance of each geotechnical property in predicting the electrical resistivities. A geotechnical property that contributes to more substantial losses in the model is assigned a higher importance score. Conversely, a feature with a score close to zero indicates the minimal impact of that feature on the predictions. According to Figure 4.12, water content exhibits the highest level of influence on the variability in the electrical resistivities. This finding also aligns with the results of Spearman's correlation analysis. The results are consistent with the literature which identifies water content as the primary factor affecting electrical resistivities (Zamanian et al. 2023b; Robinson et al. 2008). Pore water facilitates the passage of electrical current through pore spaces by moving ions, which reduces Earth's resistance (Siddiqui and Osman 2012). Shahandashti et al. (2021) showed that about 66% of the variability of electrical resistivity can be explained by the water content. Moreover, Figure 4.12 indicates that dry unit weight is the second most influencing geotechnical property affecting the electrical resistivity variations following water content. Although this finding contradicts the result of the correlation analysis, which shows a weak correlation between the dry unit weight and electrical resistivity, the existing literature shows that the dry unit weight is useful in explaining the variability in the electrical resistivities (Shahandashti et al. 2021; Alsharari et al. 2020). Changes in dry unit weights result in changes in pore spaces and interparticle contacts. Therefore, especially at low water contents, continuous pathways for the flow of electrical current can be created at high dry unit weights, which result in lower electrical resistivities (Rashid et al. 2018). The plasticity index and fine fraction demonstrate significant but

least important scores among the other geotechnical properties, implying that they have lower impacts on the electrical resistivity predictions. Lin et al. (2016) also found some correlations between the electrical resistivity of clayey soils and the plasticity index. Theoretically, fine-grained soil yields lower electrical resistivities than coarse-grained soils because they have higher specific surface areas, which promotes the transmission of electrical current (Morin 2006).



Figure 4.12 Relative importance of geotechnical properties in predicting electrical resistivity

4.2.3. Model Comparison of Deep Learning Model with ANNs, SVMs, and MLR

In this study, a deep learning model with an optimal number of hidden layers and hidden layers' neurons was adopted to assess the applicability of DNNs in investigating the non-linear and complex relationships between electrical resistivities and geotechnical properties. The performance of the trained deep learning model was then compared to artificial neural networks, support vector machines, and multivariate linear regressions – the existing methods in the literature. An ANN was trained to compare the performance of the shallow with deep network architectures in investigating the relationships between geotechnical properties and electrical resistivities. The ANN model consists of an input layer with four neurons (i.e., water content, plasticity index, dry unit weight, and fine fraction), a hidden layer with nine neurons, and an output layer with one neuron (i.e., electrical resistivity). The number of hidden layer's neurons was selected based on a trial-and-error by changing the number of neurons from one to nine and examining the model accuracy for the testing dataset through 100 iterations. An SVM with a radial basis function kernel was also trained to predict electrical resistivities based on the same geotechnical properties. The radial basis function kernel transforms the non-linear relationship between the input and output features into a linear relationship within a higher-dimensional space. The adopted deep learning model, ANN, and SVM were compared to the multiple linear regression developed by Shahandashti et al. (2021). A Box-Cox and a natural log transformation were used on the input and output variables to meet the linear regression requirements (i.e., linearity, homoskedasticity, and normality).

Figures 4.13 to 4.16 illustrate the measured and predicted electrical resistivities for the training and testing datasets by deep learning, artificial neural network, support vector machine, and multiple linear regression. The accuracy of the deep learning model (R-squared) is about 87% for the training and 70% for the testing datasets. Accordingly, it is concluded that the deep learning model yields more accurate predictions on both training and testing datasets compared to other methods. Moreover, the results imply that the deep learning model with three hidden layers is more robust than the other methods. In other words, due to the slight difference between the training and

testing accuracies for deep learning, it is concluded that the model performance remains approximately the same for predicting out-of-sample data. On the other hand, the MLR shows the lowest performance for training and testing datasets.



Figure 4.13 Training and testing accuracies for the developed deep learning to predict the electrical resistivities



Figure 4.14 Training and testing accuracies for the developed ANN to predict the electrical

resistivities



Figure 4.15 Training and testing accuracies for the developed SVM to predict the electrical resistivities



Figure 4.16 Training and testing accuracies for the developed MLR to predict the electrical resistivities

Figure 4.17 shows the testing accuracy of deep learning, ANN, SVM, and MLR for predicting electrical resistivity based on RMSE and MAE metrics. The testing accuracies show the outperformance of the deep learning model to the other models in predicting the electrical resistivities with a root mean square of errors (RMSE) of 68.9 and mean absolute error (MAE) of 17.8, followed by the ANN with an RMSE of 92.9 and MAE of 23.4. The outperformance of the deep learning model to the ability of deep network architectures to extract more meaningful information among the input and output features than shallow network architectures. Compared to all other methods, linear regression shows the most errors with an RMSE of 115.5 and an MAE of 25.7. The low performance of the multiple linear regression is because of its inability to handle the non-linear and complex relationship between electrical resistivities and geotechnical properties.



Figure 4.17 Accuracy metrics of the deep neural network, artificial neural network, support vector machine, and multiple linear regression for the testing dataset

CHAPTER 5 CONCLUSION AND FUTURE WORK

A successful design and construction of infrastructure systems such as highways and bridges highly depend on accurate estimation of geotechnical properties and understanding their spatial distributions. Insufficient and inaccurate subsurface information has a major contribution to cost overruns and delays in up to 50% of all infrastructure projects. Insufficient site investigation may also contribute to inadequate or conservative designs, leading to costly failures or increased project's costs. Hence, geophysical methods, such as electrical resistivity imaging, that can potentially transform the existing subsurface investigations are used to develop practical tools for subsurface characterization based on the data analytic approaches. This study aimed to (1) assess the presence of spatial association between electrical resistivity and geotechnical properties and propose the most appropriate spatial regression model to explain the variability of electrical resistivity values considering the spatial effects and (2) explore non-linearity and complexity of interactions between the electrical resistivity values and geotechnical properties using artificial intelligence techniques with deep structures such as deep learning model.

This research investigated the presence of spatial association between the electrical resistivity values and geotechnical properties such as gravimetric water content, dry unit weight, and plasticity index to validate the developed linear regressions in the literature. The analyses were performed based on the results of a full factorial design with 627 observations obtained from the laboratory physical property and electrical resistivity tests. Linear regression analysis was performed, and its critical assumptions were examined. Moran's I of the OLS regression residuals

showed a highly significant value, indicating that the linear regression residuals are spatially autocorrelated. Since linear regression analysis cannot consider the spatial autocorrelations among the data in the modeling, the spatial regression analysis was employed. The results showed that the SEM (spatial error model) is the most appropriate model compared to standard regression and other spatial models (SAR and SDM) in explaining the spatial variability of geotechnical properties on the electrical resistivity values. These findings indicate that the inclusion of spatial autocorrelation of residuals in the regression model could improve the performance of the regression model and lead to more accurate estimates of the effects of geotechnical properties on the variability of electrical resistivity values. These findings help engineers to have a better understanding of the effects of geotechnical properties on the variability of electrical resistivity values. The proposed approach helps create and use spatial regression models for a given site to determine the spatial distribution of geotechnical properties at each point (not necessarily those sampled using conventional site investigation methods) and conduct reliability analysis accordingly.

Moreover, an analytical approach was proposed to explore the non-linear and complex relationships between the geotechnical properties (e.g., water content, dry unit weight, plasticity index, and fine fraction) and electrical resistivity values using deep learning. A deep learning model was developed based on an empirical dataset (842 observations), comprising three hidden layers with 8, 8, and 9 neurons in each hidden layer, to relate the associations between electrical resistivities and geotechnical properties. The pairwise dependence of variables was analyzed by Spearman's correlation. Besides, the relative importance of the geotechnical properties in predicting the electrical resistivities was derived from the trained model. The performance of the deep learning model was then compared to the performance of the artificial neural network,

support vector machine, and multiple linear regression. The results showed that the water content has a significant contribution to the predictions of the electrical resistivities in clayey soils. The findings of this study also confirm that the dry unit weight plays a crucial role in electrical resistivity variations which can be attributed to the pore spaces that provide pathways for the electrical current. Furthermore, the results illustrated that the electrical resistivity of clayey soils is more influenced by the percentage of fines and plasticity index than the clay fraction and liquid limit. This study found that the deep learning model provides more accurate estimates for electrical resistivity compared to all other methods, with an RMSE of 68.9 and an MAE of 17.8. This study also showed that the deep learning model yields a more robust and generalized model since there is a slight difference between the training and testing accuracies. The outperformance of the deep learning model to ANN indicates a high level of complexity among the geotechnical properties and electrical resistivities. The research's findings help better understand the variability of electrical resistivities due to changes in geotechnical properties and improve subsurface characterization using electrical resistivity imaging technology. The proposed methodology can also be used to validate findings from electrical resistivity surveys by leveraging geotechnical properties, particularly where intrusive investigation methods are prohibited. The research results can also be applied to other applications in which understanding electrical resistivity and its relationship with the geotechnical properties is essential, such as designing grounding systems (Ackerman et al. 2013).

Future Work

This research results shed light on the spatial impacts of geotechnical properties on the electrical resistivities in clayey soils and proposed an analytical framework for explaining complexity among the electrical resistivity data. The following recommendations could be incorporated into future research:

- Due to the complexity of interactions among geotechnical and geoelectrical properties, slight changes in the geotechnical properties can result in significant variations in electrical resistivity. Future experimental designs may consider smaller increments for the factor levels such as water content to enhance the accuracy of the analysis results.
- It would be beneficial to scrutinize the effects of other geotechnical properties such as clay fraction and plasticity index on the electrical resistivity by considering multiple factor levels for each in the experimental design. These factors can be adjusted by considering different combinations of commercially available soils.
- To improve the accuracy of the spatial analysis and generalization of the findings, it would be advantageous to integrate additional data from various locations across Texas and, thereby increasing the number of neighbors in the spatial weight matrix.
- The practical recommendations for developing deep learning models in this research are limited to multivariate prediction models based on experimental data. Nevertheless, continuous geotechnical data are not always available from the field sites. In future work, it is of interest to integrate the proposed approach with a technique that leverages publicly available data to determine unknown geotechnical properties that show minimal variations over time, such as fine fraction, to simplify the developed machine leaning

models. Consequently, to be practically implemented, pairwise relationships between electrical resistivity and geotechnical properties may be established.

• To enhance the model inference through a deep learning approach, it is encouraged to conduct uncertainty analysis for evaluating the model's uncertainties and incorporating them into probability-based analyses.

APPENDIX BOREHOLE LOGS

Beaumont District (October 2019)

	DGC	The University of Texas at Arlington 416 Yates Stress 76010				BC	DRI	NG	NUI	MBE PAGE	ER 	P1)F 1
	NT											
	JECT N	UMBER _ First Visit	PROJECT LOCA		Beaumont							
DATE	E STAR	TED <u>10/13/19</u> COMPLETED <u>10/13/19</u>	GROUND ELEVA				HOLE	SIZE	inch	es		
	LING C		GROUND WATER	RLEVE	LS:							
					LING							
ווייק בספי ב אסדו	ES _		AFTER DR	LLING								
	GRAPHIC LOG	MATERIAL DESCRIPTION	SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)				FINES CONTENT (%)
	-	(SC) Gray to Brown; Sandy Clay	ss	_	8-11-3 (14)	_						
		(CL) Brown to Dark Brown; Clay; Moist				1						
		(CH) Brown; Clay	SPT ST		2-2-2 (4)	-						
10 KINGIO		(CH) Brown; Clay; Moist										
SAT			SPT		1-2-2	1						
EOTECH BH COLUMNS - GINT STD US LAB.GDT - 4/3/20 01:45 - C.:USEKSWX-A0519/UNEUKIVE - UNIVERSI 1 Y OF 1E		Bottom of borehole at 11.5 feet.										

G_P1\BOREL(LC	GO	The University of Texas at Arlington 416 Yates Stress 76010				BC	DRI	NG	NU	MBE PAGE	ER 1 0	P4 0F 1
ONTRP 1/BORELOC	CLIEN PROJ DATE	NT IECT NU	J MBER First Visit TED 10/14/19 COMPLETED 10/14/19	PROJECT NAME PROJECT LOCA GROUND ELEVA	TION _	Beaumont		HOLE	SIZE	inch	ies		
BEAUM	DRILL			GROUND WATE	RLEVE	LS:							
RESULT	LOGO	SED BY	CHECKED BY	AT TIME O	F DRIL	LING							
VIIVII	NOTE	s		AFTER DR	ILLING					AT1			
ST RESULTS/SITE RESIS	o DEPTH (ft)	GRAPHIC LOG	MATERIAL DESCRIPTION	SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)				FINES CONTENT (%)
Y IMAGING/TES			(CL) Light Brown, gray; Silty Clay; Moist;	ST									
SISTIVIT	5		(CL-CH) Brown to Light Brown; Clay with sand;	ST									
OCSIRE			(CH) Brown to Dark Brown; Clay;	SPT	-	1-0-1 (1)							
S AT ARLINGTON/LAB I	10		(CH) Brown to Dark Brown; Clay; Moist	ST ST	_	2-1-3	_						
ITY OF TEXAS]	Bottom of borehole at 11.5 feet.			(4)							<u> </u>
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от - 4/3/20 (
US LAB.GI													
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)GO	The University of Texas at Arlington 416 Yates Stress 76010					BC	DRII	NG	NU	MBI PAGE	ER	P6 0F 1
	NT	/0010	PROJEC	T NAME									
PRO.	JECT NU	MBER _ First Visit	PROJEC	T LOCA		Beaumont							
	START	ED 10/15/19 COMPLETED 10/15/19	GROUNE	ELEVA	TION			HOLE	SIZE	inch	nes		
	LING CO	NTRACTOR	GROUND			LS:							
	LING ME	THOD	AT	TIME OF	F DRIL	LING							
	SED BY	CHECKED BY	ΔΤ			ING							
			AF		LLING								
	GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)				FINES CONTENT (%)
		(CL) Brown; Silty Clay;		ST	-								
ESISTIVI		(CL) Brown; Clay		ST									
DOCSIRE		(CH) Brown; Clay		SPT		2-2-4 (6)							
LINGTONILAB		(CH) Dark Brown; Clay		ST	-								
10 XI AR						1-1-2	-						
XAS	1	Detters of boostals of 44.5 feet				(3)							
1 BH COLUMNS - GINT STD US LAB.GDT - 4/3/20 01:47 - C:/USERS/MXA0516/ONEDRIVE - UNIVERSITY													

LOGO	The University of Texas at Arlington 416 Yates Stress 76010					BOI	RIN	GN	IUM	BE PAGE	R P	1
		PROJECT										
PROJECT NU	IMBER First Visit	PROJEC			Beaumont							
DATE START	ED 10/14/19 COMPLETED 10/14/19	GROUND	ELEVA				HOLE	SIZE	inch	es		
DRILLING CO	INTRACTOR	GROUND	WATER	LEVE	LS:							
DRILLING ME	THOD	AT	TIME OF	DRILI	LING							
LOGGED BY	CHECKED BY	AT			ING							
		~							ATT	ERBE	RG	
o DEPTH (ff) GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)			PLASTICITY INDEX	
-	(CL) Gray to Light gray; Slighly moist; Silty Sand with clay; seam	sand	X ss		11-9-6 (15)	_						
	(CL) Gray; Sandy silty clay;											
<u> </u>	(CL-CH) Gray to brown; Silty Clay		SPT		1-2-4	1						
_					(0)	1						
10	(CL-CH) Gray to brown; silty clay		SI	-								
_			SPT	1	2-1-2	1						
	Bottom of borehole at 11.5 feet.				(3)							T
	bottom of borenore at 11.5 reet.											

Beaumont District (December 2019)

L	.0G0	The University of Texas at Arlington 416 Yates Stress 76010				BC	RIN	IG I	NUN	/IBE	PAGE	3 R-6 = 1 0	6 A 0F 4
CL			PROJECT	NAME	Md A	sif Akhtar							
PR	OJECT	IUMBER F	PROJECT	LOCAT	ION _	Beaumont							
DA	TE STAF	COMPLETED 12/13/19	GROUND	ELEVA				HOLE	SIZE	inch	es		
DR	RILLING C	CONTRACTOR 0	GROUND	WATER	LEVE	LS:							
DR	RILLING	METHOD Wash Boring	⊥×AT⊺	TIME OF	DRIL	LING 6.50) ft						
LO	GGED B	Y CHECKED BY	AT E	END OF	DRILL	.ING							
NC	DTES W	et Rotary set @ 8 ft due to Ground Water Table found at 6.5 ft	AFT	er dri	LING								
DEPTH	(ft) GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)				FINES CONTENT (%)
		(CL) Silty Clay; Dark Brown; Moist		ST	88								
- 5		(CH) Clay; Dark Brown to Light Brown; Moist		ST	89								
GPJ				SPT		1-1-1 (2)							
		≚ (CH) Very soft clay; Light brown to light grey; Moist		ST	89	(_)							
		(CH) Very soft clay; Light brown to light grey; Moist		ST	79								
PROJECTS		(CL) Slightly Silty Clay; Light brown with some orange and g Moist	gray;	SPT		6-8-6 (14)							
	5	(CL) Slightly Silty Clay; Light brown with some orange and g Moist	gray;	ST	58		_						
	-			SPT		4-5-9 (14)							
DITENAL 2		(SW-SC) Clayey Sand; Reddish-brown; Moist		ST	46								
29 - C:\USEF	_			SPT		7-12-9 (21)							
- 12/17/19 15:	5	(CL) Silty Clay; Light brown with some orange; Moist and So	oft	ST	88								
S LAB.GDT	-			SPT		6-8-8 (16)	-						
S-GINT STD U		(CL) Silty Clay; Light brown with some orange; Moist and So	oft	ST	100								
COLUMNS				SPT		6-6-8 (14)	-						
GEOTECH BH	5	(CL) Slightly Silty Clay; Light brown with some orange; Mois	st	ST	92								

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	200	The University of Texas at Arlington				BO	RIN	IG I	NUN	/IBE		BR-6	5 A
	JGU	416 Yates Stress 76010									TAGE		1 4
CLIE	NT SW	JIS Lab	PROJECT	NAME	Md A	sif Akhtar							
PRO	JECT N	JMBER	PROJECT I	LOCAT	'ION _	Beaumont							
DEPTH	GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)				FINES CONTENT (%)
		(CH) Clay; Gray; Moist (continued)		SPT		8-8-6 (14)							
				ST	90								
			X	SPT	-	6-6-6 (12)							
				ST	90								
Page 1	-		X	SPT		6-35-35/0"							
		Very Little Recovert; Sample looks like just cutting	×	SS									
	-		2	¶ <u>S</u> PT		8-35-35/0"							
95	-	Large Grain Sand with Rock; Gray into Whte	×	ss									
	-			SPT		8-35-35/0")							
100		Large Grain Sand; Gray: Wet	×	SS	-	9 25 25/0"							
				- <u>3</u> -1	1	<u>u-33-33/U"</u>							
105		(CH) Shale/ Clay; Blue Gray; Moist	X	ss									
		(CH) Shale to Slightly Silty Clay; Blue Gray; Moist	X	SPT SS		8-16-11/1"							
110 110			X	SPT		8-17-27/1"							
115	-	Sandy Clay to Silty Clay; Blue gray; Very Moist	X	ss									

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LC	DGC	The University of Texas at Arlington 416 Yates Stress 76010				BO	RIN	IG I	NUN	/IBE	PAGE	BR-(4 0	6 A F 4
	NT_SV	 /IS Lab	PROJECT NAM	ME N	Md As	sif Akhtar							
PRO	JECT N	JMBER	PROJECT LOC	CATIC	ON _E	Beaumont							
HLdad 115	GRAPHIC LOG	MATERIAL DESCRIPTION	SAMPLE TYPE		(RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)				FINES CONTENT (%)
	$\left \right $		X s	PT		8-35-35/0"							
	-	Sandy/ Silty Clay; Gray; Moist	X s	ss									
120		Bottom of borehole at 120.0 feet.	X s	PT		8-35-35/0"							
GEOLECH BH COLUMNS - GINT STD US LAB.GDT - 12/17/19 15:29 - C:USEKSIYUBLCUDOCUMENTSBENTLEY/GINTPROJECTSIGINT STD US LAB.GFJ													

LC	OGC	The University of Texas at Arlington 416 Yates Stress 76010				BOF	RING	G N	UMI	BEF	PAGE	R-1 ≣ 1 0	0A DF 3
CLIE	NT		PROJEC	T NAME	Beau	mont Site	Investi	gation	with E	Electric	cal Re	sisitivi	ity
PRO.	JECT N	UMBER	PROJEC	T LOCAT		Beaumont	(SH 9	S and	SH69	Inters	ection		
DATE	STAR	TED 12/12/19 COMPLETED 12/13/19	GROUNE	ELEVA	TION			HOLE	SIZE	inch	es		
DRIL	LING C	ONTRACTOR	GROUNE	WATER	R LEVE	LS:							
DRIL	LING M	ETHOD Wash Boring	AT	TIME OF	F DRIL	LING							
LOG	GED BY	CHECKED BY	AT	END OF	DRILL	.ING							
NOTE	S		AF	ter dri	LLING								
DEPTH (ft)	RAPHIC LOG	MATERIAL DESCRIPTION		IPLE TYPE UMBER	OVERY % (RQD)	BLOW OUNTS I VALUE)	CKET PEN. (tsf)	(pcf)	DISTURE NTENT (%)				S CONTENT (%)
	G			SAN	REC	υĘ	РОС	DR	₹õ	32	E L	IN	INE
		Dark Brown; Sandy Clay; Organics					1.0						
	듣글	Dark Brown: Clay					1.0						
	1						1.0						
5 6						0-0-1	0						
SLAB.						(1)	0						
	-						5						
10						0-0-1	с. С.						
						(1)	-						
N 15						4-8-8	1.0						
						(10)							
		Gray to Brown; Sandy Clay											
						5-7-7 (14)	2.0						
	11911.												
25		Gray, Brown, Tan; Sandy Clay with White seams				3-5-5 (10)							
LAB.GDI													
S0													
30						5-5-5							
						(10)	2.0						
35													

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			76010											
	PROJ	IT ECT NI	JMBER	PROJEC PROJEC	F NAME	Beau	imont Site I Beaumont	Investi (SH 9	gation 6 and	<u>with E</u> SH69	Electric Interse	cal Re ection	sisitivit 1	y
┢											AT	FERBE	RG	۲,
	(1) (1) 75	GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN (tsf)	DRY UNIT WT (pcf)	MOISTURE CONTENT (%)	LIQUID			FINES CONTEN (%)
							3-12-30 (42)	.5						
	 80		Gray; Sand Clay with wood in Bedded Organics		_		4.3.4							
ŀ	· -						(7)							
	85		Tan Gray sand with wood organics in bed		_									
AB.GPJ							12-19- 36/5"	-						
ID USL	-													
GINT S														
SINT/PROJECTSV	<u> </u>						5-6-23 (29)	-						
NTS/BENTLEY/G	95						12-50-							
3LIC/DOCUME	· _													
RS/PUE	100				_		12-50-							
- C:\USE			Bottom of borehole at 100.4 feet.				50/0"							
EOTECH BH COLUMNS - GINT STD US LAB.GDT - 12/18/19 12:18 -														

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Corpus Christi District (February 2020)

LOG	The University of Texas at Arlington 418 Yates Stress 78010				во	RIN	GN	IUM	IBE	R B PAG	R 2	01 DF 2
LIENT		PROJ	ECT NAME	The U	niversity of	Texas	at Arlin	gton				
ROJECT	NUMBER	PROJ	ECT LOCATIO	N _	Corpus Chri	sti						
DATE STA	ARTED 2/25/20 COMPLETED	2/25/20 GROU	IND ELEVATI	ON _	0 ft	0	HOLE	SIZE	inche	15		
RILLING	CONTRACTOR	GROU	IND WATER L	EVEL	S :							
RILLING	METHOD Wash Boring	¥	AT TIME OF	RILL	ING 15.0	0 ft / E	lev -15	.00 ft				
OGGED	BY CHECKED BY		AT END OF D	RILLI	NG							
OTES			AFTER DRILL	ING	222							_
- <u></u>			I YPE	37%	s Ser	PEN.	WT.	ы (%)	AT	LIMITS	RG ≻	TENT
GRAPH	MATERIAL DESCRIP	TION	SAMPLE 1 NUMBE	RECOVEF (ROD	BLOW COUNT	POCKET (tsf)	DRY UNIT (pd)	MOISTU	LIQUID	PLASTIC	PLASTICIT INDEX	FINES CON
	LEAN CLAY WITH SAND, (CL) ;Dark Brow	(D);	X ss	-	4-8-5 (11)			1. A	3	*	9	
	SILTY SAND, (SM) ;Brown; moist;		X ss		4-9-11 (20)							
	LEAN CLAY WITH SAND, (CL) ;Brown; sli	ghtly moist; ≤oft;	SPT		8-7-6 (13)							
10	SILTY SAND, (SM) ;Very Pale brown; soft;	; wet;	SPT SS		1-5-5 (10) 2-2-3 (5)	-						
15	¥		SPT		2-4-8 (12)	-						
20	SILTY SAND, (SM) ;Very Pale Brown; Wet	1	SPT	- 1995 - 2010 -	4-13-14 (27)	- 						
25	WELL GRADED SAND WITH SILT, (SW-SM	vl) ;Very Pale brown	SPT		12-33- 50/2"	8 10						
30	WELL GRADED SAND, (SW) :Light Grayis	sh;	SPT		12-33- 50/3"	-						
25	CLAYEY SAND, (SC) ;Light Gray; Stiff;		ST			4.5+						

96

60	The University of Texas at Arlington 416 Yates Stress 76010			BO	RIN	GN	IUN	IBE	R B PAG	R 2	01 0F 2
		PROJECT NAME	The U	niversity of	Texas	at Arlin	gton				
	IBER	PROJECT LOCAT	ION _	Corpus Chri	sti	_					
FOG	MATERIAL DESCRIPTION	SAMPLE TYPE NUMBER	RECOVERY % (ROD)	BLOW COUNTS (N VALUE)	POCKET PEN. (1st)	DRY UNIT WT. (pd)	MOISTURE CONTENT (%)		TERBES LIMIT	PLASTICITY 8	FINES CONTENT
		X SPT	1	12-50- 50/0"							
		ST			4.5+						
		SPT		12-34- 50/0"							
	LEAN CLAY WITH SAND, (CL) :Slight Gray; very Stiff with orangeish Stain;	ST			4.5+						
		SPT		8-33-34/2"							
	LEAN CLAY WITH SAND, (CL) ;Light Gray, Pale Brown; Stiff;	ST			2.5						
		SPT		12-50- 50/0"							
	WELL GRADED SAND, (SW) ;Pale Brown; Wet;	SPT		12-50- 50/0"							
	LEAN CLAY, (CL) ;Brown, yellowish strain; very stiff;	ST			4.5+						
		SPT		8-15-15/4"							
		ST			4.5+						
		SPT		8-17-20/5"							
	LEAN CLAY WITH SAND, (CL) ;Light Gray;	ST			4.5+						
		NUMBER	NUMBER PROJECT NAME NUMBER PROJECT LOCAT MATERIAL DESCRIPTION PROJECT LOCAT MATERIAL DESCRIPTION SPT SPT ST LEAN CLAY WITH SAND, (CL) :Slight Gray; very Stiff with orangeish Stain; ST LEAN CLAY WITH SAND, (CL) :Light Gray; very Stiff with orangeish Stain; ST LEAN CLAY WITH SAND, (CL) :Light Gray, Pale Brown; Stiff; ST LEAN CLAY WITH SAND, (CL) :Light Gray, Pale Brown; Stiff; ST LEAN CLAY, (CL) :Brown, yellowish strain; very stiff; ST LEAN CLAY, (CL) :Brown, yellowish strain; very stiff; ST SPT ST	NUMBER PROJECT NAME The L NUMBER PROJECT LOCATION 9 MATERIAL DESCRIPTION 9 9 Image: State in the interval of th	PROJECT NAME The University of Corpus Other NUMBER PROJECT LOCATION Corpus Other MATERIAL DESCRIPTION If USB (000) WW WW	PROJECT NAME The University of Texas NUMBER PROJECT LOCATION Copus Christ 01 MATERIAL DESCRIPTION UP BB WW VIII 1000 02 MATERIAL DESCRIPTION VIII 1000 VIII 1000 03 MATERIAL DESCRIPTION VIII 1000 VIII 1000 04 UP BB WW VIII 1000 VIII 1000 05 VIII 1000 VIII 1000 VIII 1000 03 MATERIAL DESCRIPTION VIII 1000 VIII 1000 04 UP BB WW VIII 1000 VIII 1000 0500 VIII 1000 VIII 1000 VIII 1000 0500 VIIII 1000 VIIII 1000 VIIII 1000 0500 VIIII 1000 VIIII 1000 VIIII 1000 0500 VIIII 10000 VIIII 10000 VIIII	PROJECT NAME The University of Texas at Aris NUMBER PROJECT LOCATION Copus Christ 01 MATERIAL DESCRIPTION UL UNIVERSITY of Texas at Aris 02 MATERIAL DESCRIPTION UL UNIVERSITY of Texas at Aris 03 MATERIAL DESCRIPTION UL UNIVERSITY of Texas at Aris 04 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 05 MATERIAL DESCRIPTION UL UNIVERSITY of Texas at Aris 06 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 07 MATERIAL DESCRIPTION UL UNIVERSITY of Texas at Aris 08 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 09 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 09 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 09 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 09 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 09 UL UNIVERSITY of Texas at Aris UL UNIVERSITY of Texas at Aris 000000000000000000000000000000000000	PROJECT NAME The University of Texas at Artington NUMBER PROJECT LOCATION Corpus Christ 01 MATERIAL DESCRIPTION 000000000000000000000000000000000000	PROJECT NAME The University of Texas at Adington NUMBER PROJECT LOCATION Corpus Christ 0 MATERIAL DESCRIPTION With 900 00 00 00 00 00 00 00 00 00 00 00 00	PROJECT NAME The University of Texas at Artigion NUMBER PROJECT LOCATION Corpus Orisis 0 MATERIAL DESCRIPTION ATTERNE 1<	PROJECT NAME The University of Texas at Adrigton NUMBER PROJECT LOCATION Corpus Christ 0 MATERIAL DESCRIPTION 0

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	MOSO		PROJECT	NAME	The U	niversity of	Texas	at Arin	gton	_		_	
E CTADT	ED 20400	COMPLETED 20100	GROUND	ELEVAT		o de	50		6170	Inch			_
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LING	THOR Wash Bar		Var	MAIER	DOUL				00.0				
CED BY	ETHOD Wash Bor	CHECKED BY		END OF	DRILL	NG	UNVE	10.410		_			_
ES			AFT	ER DRI	LING	_							
1 1							1			AT	TERBE	RG	
RAPHIC LOG		MATERIAL DESCRIPTION		APLE TYPE JUMBER	(ROD)	BLOW COUNTS	CKET PEN. (tst)	(pd)	OISTURE NTENT (%)	MIT	LIMITS NIL	DEX	S CONTEN
0	CAN CLAY M	TU CAND. (CI.) - Server Credit		WS A	RE	06	8	DR	COM	37	57	PLAS	FINE
	CEAR CEAR WI	In SAND, (CC) , brown Sanby,		X ss		4-5-5 (13)							
	LEAN CLAY WI	TH SAND, (CL) ; Brown; Slightly Moist;		X ss		9-10-8 (16)							
		1. Links Tax, Danielan, Classes, Maine, Cold.		SPT		1-3-5							
-	LEAN CLAT, (C	c) , Light ran, Rebuish, Signuy Molat, Stirr,	1	X SS		3-4-4	1						
	WELL GRADED Slightly Moist;	SAND WITH CLAY, (SW-SC) ; Light Tan; Loose	E;	ST		(8)	0.5	ĺ					
1	WELL GRADED	SAND. (SW) : Light Tan: Loose: Wet:		SPT		6-8-12 (20)							
				X ss		7-8-8 (16)							
	¥			SPT		4-8-9							
				X SS		2-4-7 (11)							
				SPT		8-11-7/2"							
				X ss		2-3-8 (9)							
				SPT		12-35-							
	WELL GRADED	SAND WITH GRAVEL, (SW) : Light Tan; Wet:		SS SS		5-12-25 (37)							
	WELL GRADED	SAND WITH SILT, (SW-SM) ; Light Gray with		SPT X SS		12-50- 50/0" 16-24-28							

LC	GO	The University of Texas at Arlington 418 Yates Stress 78010				BO	RIN	GN	UN	IBE	R B PAG	R 2	02 0F 2
CLIEN		IMBER	PROJECT N/	AME DCATI	The U	niversity of Corpus Chris	Texas sti	at Arlin	gton				_
C DEPTH	GRAPHIC LOG	MATERIAL DESCRIPTION	SAMPLE TYPE	NUMBER	RECOVERY % (ROD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pd)	MOISTURE CONTENT (%)		PLASTIC PLASTIC	PLASTICITY &	FINES CONTENT (%)
		WELL GRADED SAND, (SW) : Pale Brown, Lioght Grayish: Loose: Wet:	Å	SPT		12-50- 50/0" 5-12-17 (29)							
40		WELL GRADED SAND WITH SILT, (SW-SM) ; Pale Brown; Dense Sand;	X	SPT		8-28-21/4" 4-9-15 (24)	-						
45		LEAN CLAY WITH SAND, (CL) ; Light Grayish; Pale Brown; Sandy Clay; Stiff;		SPT		12-50- 50/0"							
50		WELL GRADED SAND. (SW) : Pale Brown, Grayish; Soft Sand;	¥	SPT SS		12-50- 50.0° 23-29-27 (56)							
55		WELL GRADED SAND WITH CLAY, (SW-SC) ; Pale Brown; Very Stiff; Slightly Moist; Seam of Clay	,X	SPT		8-20-13/3" 8-12-17 (29)							
60		LEAN CLAY, (CL) ; Pale Brown, Orange:	X	SPT		8-11-20/6"							
05			×	ST SPT		8-25-27/2"							
70		Bottom of borehole at 70.0 feet.		ST SPT		9-27-27/2							
201100 Lange													

LC	OGC	The University of Texas at Arlington 416 Yates Stress 76010				BO	RIN	GN	UM	BEI	PAC	W 2	214 OF 2
	T		PROJECT	AME	The U	niversity of	Texas	at Arlie	gton				
PROJ	ECT N	JMBER	PROJECT L	OCAT	ION	Corpus Chri	isti						
DATE	START	ED 2/25/20 COMPLETED 2/25/20	GROUND E	LEVA	TION	0 ft	1	HOLE	SIZE	inch	es		
DRILL	ING CO	ONTRACTOR	GROUND W	ATER	LEVEL	S:							
DRILL	ING M	ETHOD Wash Boring			DRILL	ING 13.0	0 ft / E	lev -13	00 ft				
LOGO	SED BY	CHECKED BY	ATEN	DOF	DRILL	ING -							
NOTE	s		AFTE	R DRI	LLING	_							
				w	2			-		AT	TERBE	RG	Ę
DEPTH	GRAPHIC	MATERIAL DESCRIPTION		SAMPLE TYP NUMBER	RECOVERY 9 (ROD)	BLOW COUNTS (N VALUE)	POCKET PEN (tsf)	DRY UNIT WI	MOISTURE CONTENT (%	LIQUID	PLASTIC	PLASTICITY INDEX	INES CONTER
		LEAN CLAY WITH SAND, (CL) ; Dark Brown sandy Clay;	X	SS		3-3-3 (6)					\square		-
		CLAYEY SAND, (SC) ; Sandy Clay with brown sand;	X	SS		4-5-8 (13)							
5			X	SPT		5-4-4 (8)							
		CLAYEY SAND, (SC) ; Brown Sand with some clayey Sand;	X	SS		4-4-8 (10)							
10		₽ WELL GRADED SAND, (SW) ; Tan; Sand; Moist;		SPT		6-6-7 (13) 4-6-8 (14)							
15		WELL GRADED SAND, (SW) ; Tan; Sand followed by Tan Coars sand with some gravel at 17.5;		SPT		8-8-8 (12) 2-3-5							
20			Ţ			8-8-5							
		WELL GRADED SAND WITH SILT, (SW) ; Coarse sand with some gravel; Wet; Dense;	×	SS		(11) 3-4-4 (8)							
25				_									
1		LEAN CLAY WITH SAND, (CL) ; Tan; Sandy Clay;	X	SPT		8-8-8/5" 5-7-8 (13)							
30		LEAN CLAY WITH SAND, (CL) ; Tan; Sany Clay followed by Ta Sand;	n	ST		(14)							
8		WELL GRADED SAND WITH SILT, (SW) ; Tan; Sand;	X	SPT		8-9-9 (18)							
35													

(Continued Next Page)

LUGU	The University of Texas at Arlington 410 Yates Stress 76010			200		GN	UW	BEI	PAG	E 2 0	0F 2
		PROJECT NAME	_The U	Iniversity of	Texas	at Arlin	ngton				
PROJECT NU	MBER	PROJECT LOCAT	ION _	Corpus Chri	isti						_
GRAPHIC CRAPHIC LOG	MATERIAL DESCRIPTION	SAMPLE TYPE NUMBER	RECOVERY % (ROD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tst)	DRY UNIT WT. (pd)	MOISTURE CONTENT (%)			PLASTICITY &	FINES CONTENT
-	WELL GRADED SAND WITH SILT, (SW) ; Tan; sand; Wet;	SPT	-	12-26-50 (76) 6-13-18							
40	Dense;	X ss		(31)							
	LEAN CLAY WITH SAND, (CL) ; Tan; Silty Sandy Clay with san seam;	nd SPT		8-19-17 (38) 9-12-13 (25)							

LC	GG	The University of Texas at Arlington 418 Yates Stress 78010			BO	RINO	GN	UM	BEI	PAG	N 2	215 DF 2
CLIER	IT	_	PROJECT NAME	The U	niversity of	Texas	at Arlin	gton				
PROJ	ECT N	UMBER	PROJECT LOCATIO	ON _	Corpus Chri	sti						_
DATE	START	TED 2/25/20 COMPLETED 2/25/20	GROUND ELEVATI	ON	0 ft	_	HOLE	SIZE	inch	es		_
DRILL	ING C	ONTRACTOR	GROUND WATER L	EVEL	.S:							
DRILL	ING M	ETHOD Wash Boring		DRILL	ING 13.0	0 ft / E	lev -13	.00 ft	_			_
LOGO	BED BY	CHECKED BY	AT END OF D	RILL	ING							
NOTE	s		AFTER DRILL	LING	_							
			w	8			2	6	AT	TERBE	RG	Ł
DEPTH (f)	GRAPHIC LOG	MATERIAL DESCRIPTION	SAMPLE TYP NUMBER	RECOVERY 9 (RCD)	BLOW COUNTS (N VALUE)	POCKET PEN (tsf)	DRY UNIT W	MOISTURE CONTENT (%	LIQUID	PLASTIC	PLASTICITY INDEX	FINES CONTE
		LEAN CLAY WITH SAND, (CL) ; Dark Brown; Sandy Clay;	X ss		2-4-4 (8)							
		SILT WITH SAND, (ML) ; Dark Brown; Sandy Clay followed by Brown Sand;	X ss		5-8-9 (17)							
5			SPT		7-8-5 (11)							
		LEAN CLAY WITH SAND (CL) - Brown: Sandy Clay:	X ss		3-3-4 (7)							
10		control the antic (es), brown, carry only,	ST		1512121	3.5						
			SPT		9-6-7 (13) 6-7-8							
15		✓ WELL GRADED SAND, (SW) ; Tan; Sand; Wet; soft;	X ss		(13)							
			SPT		12-8-10 (18)							
			X ss		2-3-4 (7)							
20			SPT		12-11-10 (21)							
		WELL GRADED SAND, (SW) ; Coarse Sand: Wet; Dense;	X ss		4-5-8 (11)							
25			SPT .		12-50-							
			Mas		7-5-16	1						
30			M		(21)							
			SPT SS		14/5 ⁻ 3-3-4 (7)							
35												

(Continued Next Page)

-	GO	416 Yates Stress 76010									PAG		
CLIEN	IT	1	PROJECT NA	ME	The U	niversity of	Texas	at Arlin	gton				
ROJ	ECT NU	MBER	PROJECT LO	CAT		Corpus Chr	isti						
2			R		8		z	M.		AT	TERBE	RG	ENT
(E)	GRAPHIC	MATERIAL DESCRIPTION	SAMPLE TY	NUMBER	RECOVERY (RCD)	BLOW COUNTS	POCKET PE (tst)	DRY UNIT V (pcf)	MOISTUR	LIMIT	PLASTIC	PLASTICITY INDEX	TINES CONT
30		LEAN CLAY, (CL-ML) : Tan; Clay Seam;	Y	PT		12-24-	+				\vdash	-	-
100		LEAN CLAY WITH SAND, (CL) ; Slight Recovery; Moist;				16/4"	1						
		; Seam of Sand;											
40				_									
		LEAN CLAY WITH SAND, (CL) ; Sand Followed by tan clay;		SS		12-48-45 (91) 22-20-25							
			(N	-		(40)							

Fort Worth District (July 2019)





















Fort Worth District (October 2020)

LOGO University of Texas at Arlington	BORING NUMBER BH-1 PAGE 1 OF 1
CLIENT TxDOT-Forth Worth	PROJECT NAME Forth Worth Slope Stabilization
PROJECT NUMBER 1	PROJECT LOCATION US 67 & W Henderson St (South)
DATE STARTED 10/7/20 COMPLETED 10/7/20	GROUND ELEVATION HOLE SIZE inches
DRILLING CONTRACTOR	GROUND WATER LEVELS:
	AT END OF DRILLING
NOTES	AFTER DRILLING
HLdg DHdv BO 0.0 MATERIAL DESCRIPTION	SAMPLE TYPE NUMBER NUMBER BLOW CONTENT (N VALUE) (N VALUE) (N VALUE)
Brown clay	ST
Dark grey clay weathered rock	10.4 %
Light gray clay with traces of limestone	TCP 50(6")-50(3")
	AU 10.4 %
	AU TCP 41-40 (81)

	200				BORING	B NUMBER BH-2
	JGU	University of Texas at Arlington				
CUE	NT Tx	J DOT-Forth Worth	PROJEC	T NAME Forth	Worth Slope Stabiliza	tion
PRO.	JECT N	JMBER 1	PROJEC	T LOCATION	US 67 & W Henderson	n St (South)
DATE	STAR	TED 10/7/20 COMPLETED 10/7/20	GROUND	ELEVATION	HOL	E SIZEinches
DRIL	LING CO	ONTRACTOR	GROUND	WATER LEVE	ELS:	
DRIL	LING M	ETHOD Auger Drilling	AT	TIME OF DRIL	LING	
LOG	GED BY	UTA CHECKED BY UTA	AT	END OF DRILL	LING	
NOT	ES		AF	TER DRILLING		
DEPTH	GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	BLOW COUNTS (N VALUE)	MOISTURE CONTENT (%)
0.0		Dark borwn clay				
20 12-42 - C WISERSHUNDROTECTS US & MENDERSON ST (SOUTH) GPJ		Dark borwn clay Light gray clay with traces of limestone Grey weathered rock		AU TCP AU	50(1")-50(1")	5.9 % 10.2 %
GEOTECH BH COLUMNS - GINT STD US LAB GDT - 1012 1201 1202				AU	50(0.5")-50(0.5")	

LUGO University of Taxas at Artington PPOLE 1 OF 1 CUENT_ICOT_Forth Worth PROJECT NUMEER 1 PROJECT NUME Forth Worth Stope Stabilization PROJECT NUMEER 1 ORPUTED 109/20 GROUND RUEVATION US 7 & W Henderson St.(South) ARTE STATED 0_Adject TLOATION US 7 & W Henderson St.(South) GROUND WATER LEVELS: ATTIME OF DRULING							BORING	S NUMBER BH-3
PROJECT NUMBER 1 PROJECT NUMB		COC	University of Texas at	Arlington				PAGE 1 OF 1
PROJECT NUMBER 1 PROJECT LOCATION US 67 & W Herefasson St (South) DATE STATED 108/20 COMPLETED 108/20 GROUND RUTE LEVELS: DRILLING NETHOD Auger Dalling AT END OF DRILLING		NT_Tx	 DOT-Forth Worth		PROJEC	TNAME Forth	Worth Slope Stabiliza	tion
DATE STARTED 10820 COMPLETED 10920 DRILLING CONTRACTOR	PROJ	ECT N	UMBER 1		PROJEC	T LOCATION	US 67 & W Henderson	n St (South)
DRELING CONTRACTOR GROUND WATER LEVELS: DRILING METHOD Auger Drilling CHECKED BY LITA AT TIME OF DRILING	DATE	STAR	TED 10/8/20	COMPLETED 10/8/20	GROUND	ELEVATION	HOLE	SIZE inches
DRELING METHOD _LOGED BY _LITAAT THE OF DRILING	DRILL	LING C	ONTRACTOR		GROUND	WATER LEVE	LS:	
LOGGED BY UTAATERO OF DRILLING	DRILL	LING M	ETHOD Auger Drilling		AT	TIME OF DRIL	LING	
NOTES AFTER DRILLING	LOGO	GED BY	UTA	CHECKED BY UTA	AT	END OF DRILL	JNG	
Hard Direction Hard ST ST 0 Dark brown stiff clay ST ST 6.8 % 0 Dark brown stiff clay ST ST 6.8 % 10 TCP 48-50(3°) 6.8 % 6.8 % 10 TCP 50(5°)-25 10.8 % 18.6 % 115 TCP 50(4.25°)-50(4.25°) 18.4 % 18.4 % 10 Crey brown stiff clay ST ST 18.4 % 10 Crey brown stiff clay AU ST 18.4 %	NOTE	s			AF	TER DRILLING		
Dark brown stiff clay ST Grey weathered rock TCP 48-50(37) 5.9 % AU AU S0(37)-50(17) 6.8 % TCP 50(37)-50(17) 6.8 % 10.8 % TCP 50(57)-25 10.8 % 18.6 % TCP 50(4257)-50(4257) 18.4 % 18.4 % TCP 50(0.57)-50(0) 18.4 % 18.4 %	o DEPTH	GRAPHIC LOG	MAT	ERIAL DESCRIPTION		SAMPLE TYPE NUMBER	BLOW COUNTS (N VALUE)	MOISTURE CONTENT (%)
Company Grey weathered rock TCP 48-50(37) 8.9 % Light gray weathered limestone AU S0(37)-50(1') 6.6 % 10 TCP S0(37)-50(1') 6.6 % 10 TCP S0(57)-50(1') 6.8 % 10 TCP S0(57)-50(1') 6.8 % 10 TCP S0(6')-25 10.8 % 11 AU TCP S0(4.25')-50(4.25') 11 AU TCP S0(4.25')-50(4.25') 11 AU TCP S0(4.25')-50(4.25') 11 AU TCP S0(0.5')-50(0) 11 TCP			Dark brown stiff clay			ST		
Image: Solution of the solution at 25 0 foot Image: Solution of the solution at 25 0 foot			Grey weathered rock			ТСР	48-50(3")	60.04
AU AU 5 50(3')-50(1') 6.8 % 10 TCP 50(3')-50(1') 6.8 % 10 TCP 50(5')-25 10.8 % 15 AU TCP 50(5')-25 10.8 % 15 AU TCP 50(4 25')-50(4 25') 18.6 % 10 Crey brown stiff clay ST ST 18.4 % 10 Crey brown stiff clay ST ST 18.4 %	립						10 00(0)	0.9 %
Image: Second	8	111				AU AU		
S TCP 50(3')-50(1') 6.8 % TCP 50(3')-50(1') 6.8 % TCP 50(5')-25 10.8 % TCP 50(4.25')-50(4.25') 18.6 % TCP 50(4.25')-50(4.25') 18.4 % TCP 500.5')-50(0) 18.4 %	NON ST		Light gray weathered li	mestone				
5 TCP 50(3')-50(1') 6.8 % 10 TCP 50(5')-25 10.8 % 10 TCP 50(5')-25 10.8 % 11 AU TCP 50(4.25')-50(4.25') 11 TCP 50(4.25')-50(4.25') 18.4 % 11 ST ST ST 11 TCP 50(0.5')-50(0) 18.4 %	<u> Ж</u>							
10 TCP 50(3')-50(1') 10 TCP 50(5')-25 10 TCP 50(5')-25 11 TCP 50(4.25')-50(4.25'') 11 TCP 50(4.25')-50(4.25'') 11 TCP 50(4.25')-50(4.25'') 11 TCP 50(4.25')-50(4.25'') 11 TCP 50(0.5'')-50(0)	<u> </u>	ᆣᅻ						6.8 %
10 Yellow brown stiff clay TCP 50(5')-25 10 Yellow brown stiff clay 10.8 % 15 AU 50(4.25')-50(4.25'') 16 ST 50(4.25')-50(4.25'') 18.4 % ST 20 Crey brown stiff clay ST 30 Crey brown stiff clay ST 31 ST ST 32 Crey brown stiff clay ST 34 ST ST 35 Battom of boreholds at 25 0 fort	2 M					ТСР	50(3")-50(1")	
10 TCP 50(5")-25 10 Yellow brown stiff clay 10.8 % 15 AU TCP 16 S0(4.25")-50(4.25") 18.6 % 70 S0(4.25")-50(4.25") 18.4 % 18.4 % 18.4 %	287	╞╧┯┫						
10 TCP 50(5")-25 15 AU 15 AU 15 TCP 16 S0(4.25")-50(4.25") 18.6 % 18.4 % 18.4 % 18.4 %	13/1 -	╞╧┷┫						
10 TCP 50(5')-25 Yellow brown stiff clay 10.8 % 15 AU 15 50(4.25')-50(4.25') 18.6 % 7 20 Grey brown stiff clay 8 15 15 16 18.4 % 18.4 % 18.4 %	Ш С	<u> </u>						
10 TCP 50(5")-25 Yellow brown stiff clay 10.8 % 15 AU 15 AU 16 Store 15 Store 16 AU 17 Store 18.6 % 18.4 % 18.4 % 18.4 %	TPR	+						
10 TCP 50(5")-25 10 Yellow brown stiff clay 10.8 % 15 AU TCP 16 TCP 50(4.25")-50(4.25") 18.4 % 18.4 %	28	╞┿┰┫						
AU TCP 50(5")-25 10.8 % 18.6 % 15 0 15 0 15 0 16 TCP 50(4.25")-50(4.25") 18.4 % 18.4 % 18.4 % 18.4 %	<u>မ် 10</u>	╞┿┯┫						
Image: Section of borehole at 25.0 fact Yellow brown stiff clay 10.8 % 15 AU TCP 50(4.25")-50(4.25") 18.4 % 18.4 %	1986 1987	+				ТСР	50(5")-25	
15 AU 10.8 % 15 AU 50(4.25")-50(4.25") 18.6 % 18.6 % 18.4 % 18.4 %			Yellow brown stiff clay					
00 AU 10.8 % 15 AU 50(4.25")-50(4.25") 18.6 % 18.6 % 20 Grey brown stiff clay ST 20 ST ST 50(0.5")-50(0) 18.4 %	Nno -							
15 AU 15 AU 10.8 % 18.6 % 18.6 % 18.4 % 18.4 % 18.4 % 18.4 % 18.4 % 18.4 %	<u>g</u>							
AU TCP 50(4.25")-50(4.25") Image: Stress of part of part pole at 25.0 fact 18.4 %	UBL .							10.8 %
10 TCP 50(4.25")-50(4.25") 18.4 % 18.4 % 18.4 % 18.4 %	ERSI					AU		18.6 %
AU ST ST ST ST ST ST ST ST ST ST	15 0					тср	50(4.25")-50(4.25")	
Grey brown stiff clay Crey brown stiff clay 20 Crey brown stiff clay 20 Crey brown stiff clay Crey brown sti	12:45							
AU AU AU AU ST ST ST TCP 50(0.5")-50(0) 18.4 %	12/20							
AU ST ST ST ST ST ST ST ST ST ST	10							
Grey brown stiff clay 3T ST ST ST TCP 50(0.5")-50(0) 18.4 %	6					U AU		18.4 %
ST ST ST			Grey brown stiff clay					
ST ST ST ST TCP 50(0.5")-50(0) 18.4 %	g 20		,			ST		
TCP 50(0.5")-50(0) 18.4 %	NT SI					ST		
TCP 50(0.5")-50(0) 18.4 %	اق ش							
TCP 50(0.5")-50(0) 18.4 %	NN							
Image: Construction	ğ							
TCP 50(0.5")-50(0) 18.4 %	あ 天							
Bottom of borehole at 25.0 feet	비 -					ТСР	50(0.5")-50(0)	18.4 0/
	8 25		Datta	m of horebole at 25.0 fact				10.41 70











Bottom of borehole at 30.0 feet.

El Paso District

LC	OGC	The University of Texas at Arlington 416 Yates Stress 76010					во	RIN	IG N	NUN	PAGE	R B = 1 0	-4 F 2		
	NT_E-	 Paso	PROJEC	T NAME	El-Pa	iso Soil Sa	mple (Collect	tion						
PROJ	PROJECT NUMBER 1				PROJECT LOCATION										
DATE STARTED COMPLETED				GROUND ELEVATION HOLE SIZEinches											
															DRILI
DRILLING METHOD Honow Stem Auger 2				AT END OF DRILLING											
LOGO		CHECKED BT	_ ^	END OF	DRILL	.ING									
NOTE															
				w.	*		z	E.		AT	LIMITS	RG	E		
o DEPTH (ft)	GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYF NUMBER	RECOVERY (RQD)	BLOW COUNTS (N VALUE)	POCKET PE (tsf)	DRY UNIT W (pcf)	MOISTURE CONTENT (3	LIMIT	PLASTIC	PLASTICITY INDEX	FINES CONTE (%)		
				SPT		6-5-4 (9)									
				N SPT		2-3-2									
5				SPT		(5) 5-4-6 (10)									
				Ç—		2.0.2									
10				SPT		(5)									
				SPT		8-18-22 (40)									
				SPT		8-6-8 (14)									
15				_		770									
				SPT		(15)									
				SPT		4-5-5 (10)									
20	2.2.12														
		SILTY SAND, (SM) Silty Sand (SM); Brown		SPT		8-10-13 (23) 5-7-10									
		CLAYEY SAND, (SC) Clayey Sand; Brown;		A SPT		(17)									
25				SPT		8-20-16 (36)	-								
		SILTY SAND, (SM) Silty Sand (SM); Brown;		SPT		9-13-15 (28)	1								
30		CLAYEY SAND, (SC) Clayey Sand (SC); Brown;													
				SPT		8-18-35 (53)									
				SPT		13-12-9 (21)									
35		(Continued Mart Deep)													

ILE 2.GPJ	LOGO		The University of Texas at Arlington 416 Yates Stress	BORING NUMBER B-4 PAGE 2 OF 2											
See H			76010												
LE 2/B		IT <u>E</u> -	Paso	PROJECT	NAME	EI-Pa	iso Soil Sa	mple (Collect	tion					
PHER	ROJ	ECT N	UMBER 1	_ PROJECT											
TIGATIONIEL PASOIBO	H (#)	GRAPHIC LOG	MATERIAL DESCRIPTION		SAMPLE TYPE NUMBER	RECOVERY % (RQD)	BLOW COUNTS (N VALUE)	POCKET PEN. (tsf)	DRY UNIT WT. (pcf)	MOISTURE CONTENT (%)	LINIT		PLASTICITY	FINES CONTENI (%)	
INVEST	-		SILTY SAND, (SM) Silty Sand (SM); Brown;		SPT		8-21-34 (55)								
TSISITE	-		LEAN CLAY, (CL) Lean Clay with Sand; Brown;		SPT	1	12-14-11 (25)	1							
RESUL	-		CLAYEY SAND, (SC) Clayey Sand (SC); Brown;		-	1		1							
NGUTEST	40 -				SPT		8-35-35								
TY IMAG	-				SPT		22-26-20 (46)								
RESISTIVI	45														
DOCSV	-				SPT		8-29-35 (64)]							
TONLAB	-				SPT		12-23-20 (43)								
ARLING	50														
XAS AT			POORLY GRADED SAND, (SP) Poorly Graded Sand wit (SP-SM), Grayish Brown;	ith Silt;	SPT		8-25-35 (60)			· · ·					
OFTE			Bottom of borehole at 50.0 feet.					-							
ERSITY															
E-UNI															
LEDRIN															
0518101															
RSMX															
- C/US															
21 13:47															
T - 3/15(
LAB.00															
STD US															
8-GINT															
OLUMN															
문문															
GEOTE															

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